GEOLOGICAL SURVEY OF CANADA.

DESCRIPTIVE CATALOGUE

OF A COLLECTION OF THE

ECONOMIC MINERALS OF CANADA,

AND OF ITS

CRYSTALLINE ROCKS.

SENT TO THE

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DESCRIPTIVE CATALOGUE

OF

ECONOMIC MINERALS OF CANADA.

BY SIR W. E. LOGAN, F.R.S.

In this Catalogue the classification of the Minerals is wholly technical, each substance being arranged under a heading connected with some one of its more prominent applications. There is given with each material the place from which it comes, and the name of the exhibitor, the latter in Italics. Beneath these is placed a list of the objects presented by each exhibitor, and a short description of the contribution, which is always terminated with an indication of the geological formation from which the substance is derived; reference being made to its Canadian designation, and in general to the English group or system in which the formation is included. These designations are also in Italics. The headings under which the Minerals are classed, are as follows:

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1.

METALS AND THEIR ORES.

IRON.

Bog Ore.

1. Radnor Forges, Batiscan .................. A. Larue & Co., Three River

- 3 pieces of bog ore of different qualities, ready for the furnace.
- Washed bog ore, ready for the furnace.
- Slag from the smelting.
- Limestone used as flux.
- Sandstone used for furnace hearths.
- Slag from the smelting.
- Limestone used as flux.
- 5 qualities of pig iron, Nos. 1, 2, 3, 4, 5. No. 3 with a polished face.
- 1 piece of pig iron, re-cast from Nos. 2, 3, and 4, using anthracite.
- 1 railway wheel, with a piece showing chill.
- 1 section of wheel, showing chill.
- 3 nail rods.
- 2 sizes of horse-shoe nails.
- 1 piece scythe iron.
- 1 pair of railway wheels and axle, which have run 150,000 miles.

Deposits of bog iron ore, in greater or less abundance, are spread out in patches on the north side of the St. Lawrence, and between it and the foot of the Laurentide Hills, all the way from Ste. Anne des Plaines to Portneuf; a distance exceeding a hundred miles. In this area, the ore seems to be most concentrated in the neighborhood of the St. Maurice and Batiscan Rivers, and iron has been smelted in the neighborhood of Three Rivers for upwards of a century. The St. Maurice Forges were established in 1737, and continued in operation until 1858. They were supplied with ore (all the hydrated peroxide), and with charcoal from the seigniory of St. Etienne, which were leased to the Smelting Company by the Crown. In 1831, according to Bouchette, from 250 to 300 persons were employed at the establishment, which had always been celebrated for the excellence of its iron; but the ore and wood becoming exhausted, and the Radnor Forges having been erected in the seigniory of Cap de la Madeleine, on the Rivière au Lard, a tributary of the Champlain River, in a vicinity where the ore and wood are still abundant, the St. Maurice forges went out of blast. The ore with which the Radnor furnaces are supplied, is derived from the seigniories of Cap de la Madeleine and Champlain, where it occurs close to the surface, in a multitude of patches distributed over the country with a thickness of from 3 to 24 inches. It is brought to the furnaces, partly by the workmen of the Company, and partly by the various farmers on whose lands the ore occurs. The chief manufacture of the Company consists of cast iron car wheels, the price of which at the forges is 2$ cents per lb. A rolling mill has recently been erected at the establishment for the rolling of malleable iron of superior quality, such as scythe iron, the price of which is 3$ cents per lb., and nail-rod iron, the selling price being 5$ cents per lb. Limestone, as a flux for smelting the ore, is obtained from the Trenton group, at the works; and sandstone for furnace hearths at the Grès rapids, on the St. Maurice, where it used formerly to be obtained by the St. Maurice Company. This quartzose sandstone belongs to
MINERALS OF CANADA.

the Potsdam formation, part of the lowest group of the Lower Silurian series of rocks. Being in this locality of a freer texture than the same beds in other parts of the province, it has been found capable of resisting a very strong heat without injury. Blocks of from 12 to 14 inches thick, 4 feet long and 20 inches wide, do not require renewal oftener than once in two years. The ore is washed at the smelting works, to free it from soil, and it then contains between 40 and 50 per cent. of iron. The quantity used annually is between 6000 and 6000 tons, producing about 2000 tons of pig iron, and the number of workmen employed varies from 200 to 400; a great many hands being required at certain periods, to excavate and bring in the ore, and to prepare and transport the charcoal.—Alluvion.

2. Vaudreuil, County of Vaudreuil ........................ Geological Survey.

a. Specimens of ore.

A bed extending under several lots on the Cote St. Charles, in the seigniory of Vaudreuil, at the confluence of the river Ottawa and St. Lawrence. The bed is in some places from 4 to 8 feet thick, and there lies beneath it, in some parts, a thin stratum of blue phosphatic iron. This bog iron ore contains about 50 per cent. of iron, but it has never been worked.—Alluvion.


a. Specimens of ore.

An interrupted bed extending over an area of 10 or 15 square miles, near the junction of the two branches of the Riviere du Sud, county of Bellechasse. The patches are from 1 to 10 acres in superficial, and from 12 to 20 inches thick. The specimens are from the property of Capt. Morin, and the ore, which has never been worked, contains about 50 per cent. of iron.—Alluvion.

Red Hematite or Oligist Ore.

1. MacNab, lot 6, concessions C and D, ....................... Geological Survey.

a. Specimens of ore.

An unworked bed of 30 feet thick, containing by analysis about 58 per cent. of iron. The bed rests upon crystalline Laurentian limestone, and is limited at the top by a magnesian limestone belonging to the Calciferous formation of the Lower Silurian era. The exposure of the ore occurs near the Fall of the Dochart, within a quarter of a mile of the shore of Lac des Chats, an expansion of the Ottawa River.—Laurentian.

2. Sutton, lot 9, range 11 ................................. Geological Survey.

a. Specimen of ore.

A bed of 7 feet thick in chlorite slate, on the property of Mr. L. H. Smith. Different portions of the bed yield from 20 to 50 per cent. of iron.—Quebec group, Lower Silurian.


a. Specimen of ore.

A bed of 7 feet thick, occurring in chlorite slate, and presenting, where exposed, the form of an anticlinal arch, which spans a breadth of 30 feet. The ore is much mixed with chlorite, and has yielded to analysis about 25 per cent. of iron. The bed is on the property of Mr. B. Mudget.—Quebec group, Lower Silurian.
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   a. Specimen of ore.

   A bed in chlorite slate. The true thickness of the bed is 5 feet, but it presents the crown of a sharp anticlinal fold, which doubles it up and gives it an apparent breadth of 10 feet. The ore may contain about 40 per cent. of iron. This bed is on the property of Mr. Reed Sweet, and with a neighboring bed of 18 feet, was formerly quarried for ore, which was conveyed a distance of 30 or 40 miles to the town of Troy, on the south side of the province line, in Vermont, and employed to assist the smelting of the magnetic oxyd procured from the serpentine in that vicinity.—Quebec group, Lower Silurian.

   N.B.—Ores similar to those of the last three localities are exposed in a great number of places in St. Armand, Sutton, and Brome, running in a bearing N. 30° E. The exposures are distributed over a breadth of about a mile, and many of them are repetitions of the same beds, through the effect of undulations. The beds are made up of hematitic iron mixed with grains of quartz and chlorite; in some the oxyd of iron predominates, constituting a rich iron ore, while in others the earthy minerals are in excess, and the rock passes into the ordinary slates of the country. These iron ores often contain a portion of titanium, as rutile, ilmenite, or sphene; in some the peroxyd is mixed with magnetic oxyd of iron.

Magnetic Ore.

   a. Specimens of ore.

   A bed of 12 feet thick, consisting of dolomite abounding in small crystals of the magnetic oxyd of iron, which equal in many specimens about 56 per cent. of the mass, thus giving an iron ore containing about 38 per cent. of metal. The ore is on the east side of a band of dolomite, varying in thickness from 12 to 32 yards, on the west side of which there is an irregular bed of red hematite 1 foot thick. Two other bands of dolomite run parallel with the one mentioned, all in the space of 100 yards, on the property of Mr. Oramel Stutson.—Quebec group, Lower Silurian.

   a. Specimens of ore.

   A mine commonly known as the Big iron ore bed of Marmora. It appears, however, not to be a single bed, but a succession of them (one measuring 100 feet in thickness), interstratified with thin bands of crystalline limestone and talcose slate, associated with diallage rock, serpentine, and epidote. The total breadth of the mass is 8 chains, and it is interstratified between gneiss and crystalline limestone, with a dip N. W. < 25° — 50°. The ore contains between 60 and 70 per cent. of iron. Many years ago a furnace was erected at Marmora to smelt it, and iron of superior quality was manufactured. More recently, different companies have for short periods renewed smelting operations, with very satisfactory results in respect to the quality of the iron produced; but the distance of the place from a shipping port has proved a serious obstacle to success. At present the furnace is not in blast.—Laurentian.

   a. Specimen of ore.

   A bed of 200 feet thick in gneiss. It is situated on Mud Lake, a part of the Rideau Canal, and is the property of Messrs. G. Chaffey and Brothers, who mine the ore, and supply it at Kingston for 24 dollars the ton, to vessels which carry it as back freight to Cleveland, on Lake Erie; whence it finds its way to the smelting furnaces at Pittsburg on the Ohio, in the State of Pennsylvania. About 4000 tons of the ore were thus exported in 1859.—Laurentian.
a. Specimen of ore.
A bed of about 90 feet in thickness. It is surrounded by gneiss, and appears to present
the form of a dome, through the summit of which there protrudes an underlying mass of
crystalline limestone. Messrs. Forsyth & Company, smelters, of Pittsburgh, commenced
mining this ore in 1854, for the supply of their own furnaces at Pittsburg, exporting the
ore by the way of Kingston, on Lake Ontario, to which it was conveyed by the Rideau
Canal. Up to 1858 they had exported about 8000 tons of ore, but the opening of the New-
borough mine, more favorably situated in regard to the shipping port, induced them to
depend upon the latter for their supply, and no ore is now exported from Hull. The ore
contains between 60 and 70 per cent. of iron. In some parts of the bed it is mingled with
a little graphite.—Laurentian.

a. Specimen of ore.
A bed of about 10 feet thick in gneiss, on the property of Mr. Thomas Loughran.—Lau-
rentian.

a. Specimen of ore.
A bed of about 20 feet thick in gneiss, on government land.—Laurentian.

7. Madoc, lot 11, range 5. G. Seymour, Madoc.
a. Specimens of ore.
A bed of 25 feet thick in gneiss, on the property of Mr. Seymour, the exhibitor: who
formerly smelted the ore at his own furnace, making from it iron of a very fine quality.
The furnace is not now in blast. The ore is very free from sulphur, and yields to ana-
lysis about 70 per cent. of iron. The beds of rock in immediate contact with the ore are
soft, black, and very micaceous, and thin veins of a similar character appear occasionally
to cut the ore bed diagonally. Nodules of radiating crystals of actinolite are disseminated
in the ore, and yellow uranite has been found investing small cracks. The ore is a
natural magnet, displaying strong polarity.—Laurentian.

a. Specimen of ore.
A bed of about 12 feet thick in gneiss. The ore which contains between 60 and 70 per
cent. of iron is of very uniform character. The proprietor has recently mined about 800
tons, which arc about to be drawn to the Rideau Canal. A small quantity of it has been
tried at Mr. Gizowski's foundry, at Toronto, and the ore is found to be well adapted for
lining furnaces.—Laurentian.

a. Specimen of ore.
A bed in gneiss, the property of the exhibitor.—Laurentian.
Ilmenite with Rutile.

   a. Specimen of ore.
   A bed of 90 feet thick, which is exposed for 300 feet on the strike, and is traceable for about a mile. The ore has yielded to analysis:

<table>
<thead>
<tr>
<th>Component</th>
<th>Percentage</th>
</tr>
</thead>
<tbody>
<tr>
<td>Oxide of titanium</td>
<td>48.60</td>
</tr>
<tr>
<td>Protoxide of iron</td>
<td>46.44</td>
</tr>
<tr>
<td>Magnesia</td>
<td>3.60</td>
</tr>
<tr>
<td></td>
<td>100.00</td>
</tr>
</tbody>
</table>

   In some parts of the bed, rutile is disseminated in the ilmenite in small red crystalline grains. The ore is interstratified in anorthosite rock.—Laurentian.

LEAD.

Galena.

   a. Undressed ore from the lode.
   b. Hand-picked prills.
   A vein transversely cutting stratified limestone, which dips about S. W. 21' and rises northward into a hill about 700 feet in height, constituting Gaspé promontory. The vein has a width of about 18 inches, and is composed of calc spar holding disseminated masses of galena. A trial shaft of 20 feet in depth has been sunk on the vein, and from this and several small veins running parallel with the main one, about 6 tons of ore of 60 per cent. have been obtained.—Lower Helderberg group, Upper Silurian.

2. Upton, lots 50, 51, range 4 ........................ James Wright & Co.
   a. Undressed ore.
   A bed composed of dolomite, with irregularly disseminated patches of galena, varying in thickness from 1 to 4 inches, but not easily traceable on the strike. The bed occurs in the upper part of a band of dolomite of from 200 to 300 feet thick, which has been followed a long distance through the country.—Quebec group, Upper Silurian.

   a. Prill ore as taken from the lode.
   b. Hand-picked prills.
   c. Sorted ore, prepared for the crusher.
   d. Pig lead run from the furnace.
   e. Slag, from the smelting of 80 per cent. ore.
   f. A plan of the mine by Mr. E. Balfield.

   A vein cutting nearly horizontal beds of grey, geodiferous, brown-weathering dolomite. The vein is composed of calc spar, and has a breadth varying from 21 to 6 feet, in which the galena is disseminated in a width of from 8 to 24 inches. In some portions the
A vein is almost dead ground, while in others, judging by the eye, it would yield nearly 2 tons of 80 per cent. ore per fathom. The bearing of the lode is about N.W., and it underlies to the north-eastward, about a foot in a fathom. A trial shaft has been sunk on the lode to the depth of 57 feet, and the working of 75 fathoms of ground, in 1858, yielded 30 tons of ore of 80 per cent. A smelting furnace was erected to reduce the ore, and a 10 horse-power engine used to give blast to the furnace and dry the shaft, but a considerable spring of water having been struck, it became necessary to erect a more powerful engine, and one of 50 horse-power has just been completed. The dolomite is underlaid conformably by sandstone, which crops out about a mile from the mine, and is unconformably supported by crystalline limestone and gneiss of Laurentian age. About 105 fathoms south-eastward from the main shaft, a counter lode joins the main lode, at an angle of about 20°; its course being nearly N.N.E. and S.S.W. At the junction of the two lodes a shaft has been sunk in sandstone, to a depth of 21 feet, and in the excavation of the pit, in which the united lodes have a breadth of 10 feet, there have been obtained about 7 tons of ore of 20 per cent.—*Calciferous formation, Lower Silurian.*

   a. Undressed ore.
   b. Plan of lodes by Mr. E. Banfield.
   Ore from a vein cutting crystalline limestone, and running N. 60° W. The vein has a thickness of from 6 to 12 inches, and is composed of calcite, in which the galena is disseminated in lumps; which, in a trial shaft of about 50 feet, sunk in 1854, on the land of Mr. Buel, were sufficient to pay the expenses. The largest of these lumps may have been 5 or 6 inches in width. A counter lode diverges from the main one near the shaft, and in this neighborhood occur 4 additional lead-bearing veins, running parallel with the main one, all contained in a breadth of about 1000 feet. They run obliquely across the lots, and thus intersect the lands of several proprietors. On lot 4 of the same range, Messrs. Foley & Co., of Montreal, have sunk a small shaft on one of the lodes.—*Laurentian.*

   a. Undressed ore.
   Ore from one of 5 nearly parallel lodes, cutting crystalline limestone, in a breadth of about a quarter of a mile, on the property of Mr. Weston Hunt, of Quebec. The gangue of the lode is a mixture of heavy spar and calcite. About a mile to the eastward of these are other nearly parallel lodes, also cutting crystalline limestone, on land belonging to the same proprietor. Shallow trial shafts were many years ago sunk on some of these, but what quantity of lead ore was obtained in them has not been ascertained. On lot 13, range 5 of Bedford, Messrs. Foley & Co., of Montreal, have sunk a trial shaft to a depth of 14 feet, on a lead-bearing lode of 6 inches, of which the gangue is heavy spar. It cuts crystalline limestone, and reaches gneiss, and in both rocks shows good bunches of ore. This lode is about 3 miles south-west from those first mentioned, and runs parallel with them.—*Laurentian.*

N.B.—The distance between the Lansdowne and Bedford lodes is about 25 miles; they bear for one another, and it appears not at all improbable that the veins in the two localities may be identical, or belong to one group. If a line from the Bedford to the Lansdowne lodes were continued 25 miles farther, it would cross the St. Lawrence, and strike across to a well-known group of veins of lead ore intersecting Laurentian gneiss. Though just now abandoned, some of these are supposed to be still unexhausted, and two of them are known at one period to have yielded a great quantity of ore; one of them as much as $142 worth to a fathom. The Ramsey lode belongs to a series of veins which run parallel with those of Bedford, at a distance of about 40 miles to the north-eastward, and, although the two groups cut different rocks, both are probably of one age, which would be that of the *Calciferous formation* of the *Lower Silurian* series.
1. Escott, lot 7, range 2, near Brockville. **Geological Survey.**
   a. Yellow sulphuret of copper, with iron pyrites and magnetic oxyd of iron, from a bed running N. E. and S. W.
   This bed is interstratified in gneiss, and consists of magnetic oxyd of iron of about 6 inches thick, which near a cutting, made for the convenience of the Grand Trunk Railway, was ascertained to be underlaid by copper pyrites. This was mined, and found to be a lenticular mass, extending about 12 feet continuously in the bed, with a thickness of 10 inches in the middle. This mass was nearly pure copper pyrites, in which thin leaves of hydrated peroxide of iron ran in cracks and joints. In some parts calc spar was present in short, thin veins and small specks, and iron pyrites was disseminated in others, increasing in quantity as it approached the north-west side; into which the copper pyrites appeared to run for short distances. Traces of cobalt occur in the iron pyrites. About 20 tons of the copper ore were obtained, but after this became exhausted, no excavation through the dead ground was made in search of a farther quantity. It is stated, however, that another mass of copper ore has been found at the surface a short distance to the S. W., but the details relating to it have not been ascertained, farther than that it is said to be 3 feet thick, and that a sample, which was an average of 9 inches of the breadth, yielded 10 per cent. of copper to the analysis of Mr. McFarlane. **Laurentian.**

2. Bruce Mines, Lake Huron ................. **Montreal Mining Co., Montreal.**
   a. Yellow and variegated sulphurets, from the lode.
   b. " " rough dressed.
   c. " " jigged.
   d. Rough waste from jigging on copper bottom sieves.
   e. Plans of the mine.
   At the Bruce mines, a group of lodes traverses the location in a north-westward direction, intersecting a thick mass of interstratified greenstone trap. The strata here present an anticlinal form, the lodes running along the crown of it. All of the lodes contain more or less copper ore, which is disseminated in a gangue of quartz. The main lode, which is worked with another of about the same thickness, is on the average from 2 to 4 feet wide, and as indicating the generally productive character of the ground, it may be stated that in a careful examination made in 1848, about 3900 square fathoms of the lodes were computed to contain about 61 per cent. of copper. The quantity of ore obtained from the mine, since its opening in 1847, is stated to be about 9000 tons of 18 per cent. The quantity obtained in 1861 was 4224 tons of 17 per cent. The deepest working is 50 fathoms from the surface. The number of men employed is 34. Smelting furnaces, on the reverberatory principle, were erected at the mine in 1853; the fuel used in reducing the ore was pit coal imported from Cleveland; but after a trial of 3 years, the Company themselves ceased smelting, and subsequently leased their smelting works to Mr. H. R. Fletcher. At present the ores are in part sent to the Baltimore market, and in part to the United Kingdom. **Huronian.**

3. Wellington Mine, Lake Huron ............... **West Canada Mining Co.**
   a. Yellow sulphuret, from the lode.
   b. Yellow and variegated sulphurets, prills.
   c. " " jigged.
   d. " " buddled.
   e. " " crushed.
   f. Rough waste from jigging.
   g. Fine waste from tyes.
   h. Plans of the mine, by Mr. C. H. Davis.
The lodes of the Wellington mine are probably a north-westward continuation of the Bruce mine group. They are of the same general character, some of them occasionally reaching a thickness of 10 feet. They occur on the ground of the Montreal Mining Company, from whom they are leased by the West Canada Mining Company at a royalty, and continue into the adjoining sett, called the Huron Copper Bay location, where also they are worked by the West Canada Mining Company. The quantity of ore obtained by this Company, from the Wellington mine, since 1857, is a little over 6000 tons of 20 per cent. In 1861, the quantity was 1175 tons of 19 per cent., and from the Huron Copper Bay mine, probably about 1300 tons; making the total quantity obtained by the two mining companies in that year about 3000 tons. The deepest working on the West Canada Company's ground is about 20 fathoms. The number of men employed on the Wellington and Copper Bay mines is supposed to be about 250. All of the ore raised by this Company is sent to the United Kingdom.—Huronian.

4. Acton Mine, Acton, lot 32, range 3. . . . W. H. A. Davies and C. Dunkin, Montreal,

a. Variegated sulphuret, from the bed.
b. " " jigged.
c. " " rough dressed.
d. " " tye work.
e. Waste from the tyes.
f. A polished slab of the conglomerate ore.
g. Rock of the country at the mine.
h. Plan of the mine, by Messrs. Willson and Robb.

The ore of the Acton mine occurs in masses subordinate to the stratification, at the summit of a band of greyish-white and reddish-grey compact subcrystalline dolomite, from 200 to 300 feet thick, belonging to the base of the Quebec group. The dolomite is divided into massive beds; it is associated with a good deal of chert, and often encloses mammillated fibrous concretionary forms, resembling those of travertine. At the summit, the dolomite often terminates in a breccia or conglomerate, with angular and rounded masses of limestone, intermingled with ragged, irregular masses of chert. In many places the dolomite is marked by the occurrence of the yellow, variegated and vitreous sulphur-crets of copper, which present themselves in patches running with the stratification. In the neighborhood of these, many veins and strings of quartz intersect the rock in various directions, and hold portions of the sulphur-crets of copper. The copper-crets, which often contain native silver, appear to be more abundant in the upper part of the rock. At Acton, the conglomerate is separated from the main body of the dolomite by between 80 and 90 feet of dark grey or black slates, intermixed with diorite; in these the conglomerate lies in large isolated masses, running parallel with the summit of the main body of the dolomite. On the opening of the mine, the sulphur-crets, where most abundant, appeared to occupy a position immediately near some of the isolated masses of conglomerate, and partially to surround them; in some parts constituting the paste of the conglomerate. As the work proceeded, many slips and dislocations, of no great magnitude, were found to cut the strata. Some of them appear to run with the strike, and others in two or parallel series, oblique to one another. These disturb the regular continuity of the copper-bearing bed, producing apparent undulations in the dip, and causing the diorite or the limestone to protrude into the copper ore, or unexpectedly to interrupt one another. The ores were found to be concentrated in three large bunches, occurring in a length of about 120 fathoms. Proceeding south-westwardly, the space occupied by the most northern bunch, from a breadth of a few inches, gradually widened out to about 10 fathoms, in a length of about 40 fathoms; beyond which it appeared to be thrown about 14 fathoms, obliquely to the westward. The general bearing of the succeeding two bunches was still to the south-west. They were about 15 fathoms apart, and the larger or more southward one
swelled to a breadth of more than 15 fathoms. The depth to which the ground has been worked on the general slope of the bed, is about 10 fathoms; the cupferiferous rock at this depth has a breadth of about 12 feet in a shaft on the northern bunch, and shows rich ore in the floor and the parts adjacent; but with the exception of what is called Pike's pit, in the most southern part, the floors of the other bunches do not at present appear to show that same abundance of ore which characterized the upper part. The working of the mine, however, up to the present time, has been confined to the extraction of the rich ore which was in sight. Little or nothing has been done for discovery, and it cannot be said how near to the present floor of the mine there may be other bunches similar to those that have been excavated. Beyond these bunches, in opposite directions on the surface, the ore becomes more scattered in the strata; but there is evidence of its continuance for several hundred feet, in spots and patches, occasionally aggregated into bunches of much less importance than the three principal ones. In the first few weeks' work in 1859, about 300 tons of ore, containing nearly 30 per cent. of copper were quarried in open cuttings from two of the masses, without making much apparent impression on the quantity in sight. The total quantity sent from the mine up to the end of 1861, is said to be nearly 6000 tons; holding on the average about 17 per cent. of copper.—Quebec group, Lower Silurian.

   a. Yellow sulphuret, from a bed.

   The band of dolomite, which sinks with a north-westward dip at Acton, rises again at Upton, on the opposite side of a synclinal form, at a distance of about six miles. Here, about 20 feet in the upper portion of the band are marked by the yellow sulphuret of copper; which is disseminated in the rock, as if in a bed, the ore being most abundant in the lower part of the bed. The rock is at the same time cut by many reticulating strings and veins of calcspar, which hold ore. An open cutting has been made on the outcrop of the bed; the quantity of ore obtained, is stated by the proprietors to be 40 tons, and a sample, represented by them to be an average one, has yielded to the analysis of Mr. C. Robb 14 per cent. of copper. The quantity of rock which has been excavated is uncertain.—Quebec group, Lower Silurian.

   a. Yellow sulphuret, from a bed.

   From the position where the rock has been wrought in the previous mine, the band of dolomite runs south-westward for nearly a mile, and then appears to be thrown upwards of a mile to the south-westward, by a dislocation. Bissonette's mine is on the south-west side of the dislocation, and apparently in the same stratigraphical place in the band, as the Upton mine. The bed is about 20 feet thick, and the ore lies in disseminated masses of various sizes, up to 50 inches long, by from 6 to 9 inches thick. The bed might probably yield from 1 to 2 of a ton of 10 per cent. to a fathom.—Quebec group, Lower Silurian.

   a. Yellow, variegated and vitreous sulphures, from a bed.

   This ore occurs in masses, disseminated in what appears to be a bed, of uncertain thickness, in the same band of dolomite as that of the Acton mine. An experimental shaft has recently been sunk on it to a depth of about 5 fathoms, in which good bunches of ore have been met with. About 4 tons of 30 per cent. ore have been obtained from the excavation.—Quebec group, Lower Silurian.

a. Yellow sulphuret, from a lode.

At this mine, several veins, carrying more or less copper, intersect a mass of magnesian limestone, which is supposed to belong to the same band as that of the Acton mine. The veins have a general bearing north-eastward, and trial shafts have been sunk on three of them, the thicknesses of which vary from 2 to 12 inches. The vein-stone is calc spar, with a little quartz, occasionally mixed with portions of the wall rock. On the most north-western vein, the excavation is 2 fathoms deep, and reaches black shale beneath the limestone. On the middle one, which is 18 feet to the south-east, the excavation is 4 fathoms deep, again reaching black shale. On the third, 24 feet farther to the south-eastward, a shaft sunk about 4 fathoms, is still in magnesian limestone. In this shaft, the vein underlies to the south-eastward, about a foot in a fathom, and in a breadth of from 6 to 12 inches, shows good bunches of ore, mixed with calc spar and wall rock.—Quebec group, Lower Silurian.

9. Black River Mine, St. Flaviën ............... Shaw, Bignol & Hunt, Quebec.

a. Yellow sulphuret, from a bed.

At St. Flaviën, about 5 leagues above the Chaudière, and two leagues from the St. Lawrence, red shales occur, underlaid by a band of amygdaloidal diorite; this appears to occupy the place of the magnesian limestone, to which the band at Acton belongs. It is between a quarter and half a mile wide, and limestones occur both at the summit and at the base of the band, which in these parts appear to be of a concretionary, or conglomerate and brecciated character, being composed, particularly at the base, of rounded and angular masses of amygdaloidal diorite, varying in diameter from 2 inches to 2 feet. Many of these are calcareous, and much of the rock is red. The interstices among the masses are filled with calc spar, which is transversely fibrous towards the walls, and incloses crystallized quartz in the centre. This band is highly cupriferous, and ores of copper occur both in the beds, and in veins or lodes which cut them: the bearing of the veins, however, being with the strike. The ore in the beds is copper pyrite, large masses of which, similar to the one exhibited, occur, associated with the limestones at the top. The veins, in addition to copper pyrites, hold the varigatcd and vitreous sulphur-ets. In one spot, native copper occurs in small masses in the conglomerate at the base of the diorite. The whole band has a striking resemblance to some of the rocks of the Upper copper-bearing series of Lake Superior.—Quebec group, Lower Silurian.

N.B.—A band of diorite very similar to the one above mentioned, and perhaps a continuation off, occurs at Drummondville on the St. Francis, where the band is a mile wide. On lot 1, range 1 of Wodover, it holds yellow, variegated and vitreous sulphur-ets of copper, which run in 6 or 7 thin veins or courses, formed by breaks and slips in the diorite, within a breadth of 260 yards.

The rocks of the Quebec group, which are almost wholly on the south side of the St. Lawrence, are distributed in long narrow parallel synclinical forms, running N. E. and S. W. For the convenience of description, these in the geology of Canada, have been divided into: 1st, The Lauzon and Farnham synclinal, which is the one most to the N. W.; 2nd, The Shipton and St. Armand synclinal, continued to the K. E., in the Shipton and Leeds synclinal. Between these two synclinals runs the Bayers and Stanbridge anticinal, and beyond them, to the S. E., is the Danville and Sutton anticinal. From this there branch, in the neighborhood of the St. Francis, the Sutton Mountain anticinal, and the Melbourne and Patton anticinal. The six copper-bearing beds and veins that have been mentioned, 4—9, are all included in the Lauzon and Farnham synclinal.
   a. Variegated and vitreous sulphures, from Hall's lode.
   b. " " " from Campbell's lode in Kent's shaft, at 30 fathoms.
   c. Yellow, variegated, and vitreous sulphures, from lowest bed.
   d. " " " from highest bed.
   e. Variegated and vitreous sulphures, dressed on copper-bottom sieves.
   f. A plan of the mine, by Mr. Herbert Williams.

At Harvey's Hill mine, there occur in a breadth of about 1000 feet, 8 courses with a north-easterward bearing, composed chiefly of quartz, with various proportions of biter spar, chlorite and calcspar. They all cut the strata, with an underlie, at high angles, to the north-westward, and hold, in greater or less quantities, the yellow, variegated and vitreous sulphures of copper. These quartz courses which appear to have lenticular forms, occasionally extend upwards of 100 fathoms horizontally; some of them have shown a width of as much as 7 feet in the thickest part, occasionally carrying, for short distances, as much as 2 tons of 20 per cent. ore to a fathom. The rock of the country is a talcoid mica slate, which from its lustre is called nacreous slate. To prove the quartz courses in a downward direction, an adit level is being driven through these slates, from the north side of the hill, at a level of 37 fathoms below its summit. The length of this adit, when complete, will be 220 fathoms. The same sulphures of copper which characterize the quartz courses, occur also in beds conformable with the stratification. Of these there are 3 at Harvey's Hill. The lowest one, resting on a 6 feet bed of soapstone, is 6 inches thick; 15 fathoms above this there is another, 3 inches thick, and 20 fathoms, still higher, one varying in thickness from 20 to 30 inches. In these beds the ore is distributed in the nacreous slate in patches, generally of a lenticular form; they are usually thin, but sometimes attain \( \frac{1}{3} \) to \( \frac{1}{5} \) of an inch in the thickest part, and occasionally present, in section, lines of 6 inches, or even 12 inches in length. The patches interlock, one overlapping another, with variable distances between; while many single crystals and grains of ore are disseminated through the whole thickness of the beds. The quantity of ore obtained from the mine is uncertain; the number of men employed is about 50.—Quebec group, Lower Silurian.

   a. Yellow sulphuret, from a vein.

The ore is disseminated in a vein, slightly oblique to the stratification of a quartz-chloritic rock, frequently studded with nodules of orthoclase feldspar, often surrounding small centres of quartz; the nodules give to the rock the aspect of an amygdaloid trap. The bed is about 18 inches wide, and the rock is supposed to occupy a higher stratigraphical place than the Acton dolomite. The vein is traced, running N. E., for 90 fathoms. Two small excavations, each of a few fathoms in length, have been made in the outcrop, to the depth of 2 feet, and in these the variegated and vitreous ores are mixed with the yellow sulphuret.—Quebec group, Lower Silurian.

   a. Variegated and vitreous sulphures, from a bed.

The bed to which this ore is subordinate, is of the same character as that of the St. Francis mine. It dips north-westward, at a high angle, and is about 12 inches thick; a shaft has been sunk in it to a depth of 3 fathoms. 10 fathoms to the east and 15 fathoms to the west of this, other copper-bearing beds occur, composed of an amygdaloidal chloritic rock like that of the St. Francis mine, one of them 3 feet and the other 6 feet thick. In these the ore is sparingly disseminated.—Quebec group, Lower Silurian.
13. Coldspring Mine, Melbourne, lot 6, range 2.---Flowers, Mackie & Co.

a. Variegated and vitreous sulphurets, from a bed.

b. Plan of the mine, by Messrs. Wileen and Robb.

The bed from which these specimens are derived, is composed of quartz and nacreous slate, in which the ore is disseminated in thin interlocking lenticular patches, and in grains; as in the beds of Harvey's Hill mine. The dip of the strata is north-westward, at an angle of about 45 degrees. Last summer, a shaft was sunk to cut the bed at 7 fathoms, but none of the ore has yet been stoped. In a breadth of 120 feet across the strata, on one side of the shaft, and 80 feet on the other, there are several parallel bands of cupriferous strata, marked chiefly by the green carbonate of copper, but showing occasional indications of the variegated and vitreous sulphurets. What the productiveness of the ground may be, has, however, not been ascertained.—Quebec group, Lower Silurian.


a. Variegated and vitreous sulphurets, from a bed.

b. Plan of the mine, by Mr. J. Richardson.

c. The ore occurs in nacreous slate, in which it is disseminated in thin, lenticular patches and in grains, as in Harvey's Hill mine. The thickness varies from 1 to about 4 feet, and the bed dips N. W. < 86°-90°. In this attitude it is visible for 170 yards, and is traceable for a mile, running parallel with a band of dolomite, which is removed from it about ¼ a mile across the strike, to the eastward. Nodules of magnesian limestone are disseminated in the slate, close along the east side of the part charged with copper ore. The band of dolomite is supposed to be in the same stratigraphical place as that of Acton, but it occurs on the eastward side of a distinct synclinal form, the axis of which is separated from that to which the Acton band belongs, by about 12 or 15 miles. A sample of the whole breadth of the bed, where it is 4½ feet, yielded to analysis 4½ per cent. of copper. A pit of 10 fathoms deep was, last year, sunk down the incline of the bed, and a small quantity of the ore stoped out at the bottom.—Quebec group, Lower Silurian.

15. Craig's Range Mine, Chester, lot 8, range 5.---G. D. Robertson & Co.

a. Vitreous sulphuret, with green carbonate, from a vein.

The vein, which is composed of quartz, has a thickness of about 2 feet. It runs with the strike, in chloritic slate, and has been uncovered for a fathom or two along it. It shows enclosed masses of the ore, but the work done is not sufficient to authorize any statement in regard to the quantity.—Quebec group, Lower Silurian.

N.B.—The 6 copper-bearing beds and lodes, 10-15, are all within the Shipton and St. Armand synclinal. Indications of copper occur in a great number of localities in this synclinal, in testing a good many of which there have been expended considerable sums. These indications run through Stukely, Ely, Melbourne, Cleveland, Shipton, Chester, Halifax, Inverness and Leeds, and cross the Chaudiere into the seigniory of St. Mary.


a. Yellow and variegated sulphurates, from a bed.

The ore of this mine occurs at the summit of a band of slaty dolomite about 300 feet thick. At the spot the rock dips S. 10° E. < 46°, and runs thence in a general eastwardly direction. For a thickness of about 30 feet, in which nacreous slate is mixed with the dolomite, the ore is disseminated in lenticular patches of various sizes, sometimes measuring several feet in length, with the thickness of an inch or more in the centre. The patches interlock among one another, and appear to be in sufficient abundance to
promise a profitable return. The dolomite crosses the north branch of the Nicolet River, producing a considerable fall in the stream, which is 20 feet wide, and would afford abundant water power for crushing and dressing the ore.—Quebec group, Lower Silurian.

   a. Iron and copper pyrites, from a bed.
   This appears to be a large mass of iron and copper pyrites, subordinate to the strata, which here consist of calcareous serpentine, and run N. E. and S. W., with a dip about S. E. < 50°. The entire thickness of the mass is uncertain, but the breadth in which the sulphuriferous beds are more or less mingled with the rock is probably not less than 20 feet. In some parts, sulphuriferous iron prevails almost to the exclusion of that of copper, while in others there is as much as 8 per cent. of copper; some parts assume the aspect of what, among Cornish miners, is termed bell-metal ore. An opening has been made in the mass, 8 feet in length, 4 feet in height, and 4 feet wide; in this, the two sulphuriferous beds are unequally mixed with one another, but nearly free from the rock of the country.—Quebec group, Lower Silurian.

   a. Yellow sulphur, from a bed.
   This bed is 5 feet thick, and occurs in a calcareous chloritic slate. The mine has been opened on a twist in the stratification, giving three courses of ore in the breadth of 80 feet, but the general plane of the bed dips about S. < 65°. A pit has been sunk on the incline of the bed to a depth of about 6 fathoms from the surface, and the ore obtained from the excavation, without any dressing, has been sent to Boston, where it has yielded on an average about 8 per cent. of pure copper. The quantity of such ore obtained from the bed by 5 men in 5 months, is about 100 tons. The bed is traceable for a considerable distance in opposite directions from the pit, and carries copper as far as it has been tried. The horizon of the strata of this mine is supposed to be higher than the dolomite of Acton, and to be approximately equivalent to the chloritic slates of the Shipton and St. Armand synclinals. The rock of Haskell Hill composes a belt of ridgy land, running from Owl's Head to Ham Mountain, forming in its progress the Stodd's mountains. It spreads out to a width of about 7 miles on the St. Francis, and shows indications of copper near Sherbrooke, on the land of Mr. Sheriff Bowell, and in several other places. A vein on lot 17, range 7, of Ascot, within a mile of Sherbrooke, in addition to the yellow sulphur of copper, has been found to hold traces of gold.—Quebec group, Lower Silurian.

N.B.—Besides the 15 described localities, 4-18, of copper-bearing beds and veins belonging to the Quebec group, nearly 200 additional localities, on separate lots of 200 acres each, are known, on which indications of the metal occur.

Native.

   a. Mass from a lode.
   On the Canal Earth, at the east end of St. Ignace Island, the vein from which the above specimen is derived, cuts a thick mass of amygdaloidal diorite, which lies conformably with the strata, there dipping S. < 9°. The vein is about 4 or 5 inches wide, and holds masses of native copper, many of them weighing upwards of 100 lbs., accompanied by native silver, in a gauze of calc spar. The underlie of the vein is N. < 70°. About 47
MINERALS OF CANADA.

17

3 feet south from this vein, there is another about 12 inches wide, of which the vein stone is calcite, with a little quartz, associated with fragments of wall rock. Among these substances are disseminated masses of vitreous copper, accompanied with native silver. The underlie of the vein is N. < 60°. These two veins would meet downwards about 25 fathoms from the surface, and with a view of testing them, the Montreal Mining Co., to whom the location belongs, in 1846 commenced the sinking of a shaft, about 34 feet north of the native-copper vein. It was carried down 10 fathoms; a drift from it southward, then intersected the native copper vein at a distance of about 20 feet, thus proving the continuance of this vein for a depth of 10 fathoms; but the Mining Company, having at that time determined to concentrate all their energy upon the working of the Bruce mine, the experiment was carried no farther. These veins, variously modified, can be traced to the westward for 1 mile, along the whole length of St. Ignace Island, carrying native copper and native silver, with the vitreous sulphuret of copper, in greater or less quantity the whole way; and also to the eastward across the northern part of Simpson's Island.—Quebec group, Lower Silurian.


a. Nodules, from a bed.

b. Gangue, or rock of the bed.

c. Plan of the Quebec mine by Mr. D. S. Cutting.

On the north side of Michipicoten Island there is a considerable mass of greenstone, several bands of which are of an amygdaloidal character, and some of them are associated with beds of sandstone. Towards the west end of the island these rocks present a low surface for about 400 or 500 feet, and then rise into a cliff of 200 or 300 feet. In the cliff, the greenstone is marked by crystals of analcime and quartz, occurring in druses. The whole mass lies conformably with the strata, dipping south-eastward. Native copper, associated with native silver, is disseminated in several parts of the mass, and these more particularly characterise an amygdaloidal bed, 2 feet thick, which is underlaid by a band of sandstone, and has been mined to a small extent by the Quebec Mining Company. In this bed, the copper is distributed in irregular nodular masses of various sizes, from grains no larger than snipe shot, to fantastic forms of 5 and 6 inches in diameter; the quantity of metal in the bed being according to Mr. J. L. Wilson, equal to about 5 per cent. Small nodules of calcite occur with those of copper. About 7 miles to the north-eastward, the bed is cut by a vein, in which copper and silver appear to be associated with ores of nickel, in the forms of a silicate of nickel, containing 35 per cent. of the metal, and of a mixture of the arseniuret of nickel and copper, containing between 17 and 27 per cent. of nickel. These ores were detected by Mr. Storpy Hunt in the refuse thrown aside in a crop trial made on the bed, by Mr. Bonner, in 1854, and it is said that a considerable quantity of the silicate was thrown into the lake, after being stamped and washed, for the purpose of extracting from it the native silver.—Quebec group, Lower Silurian.


a. 450 lbs. in a single sheet, from a vein.

The promontory of Mamainse consists of various layers of coarse conglomerate, and of greenstone, much of which is of an amygdaloidal character. According to the description of Dr. Dawson, one of the bands of greenstone is intersected by a narrow fissure, running nearly in the strike of the beds, or north and south. Its greatest width is about 6 inches, and in some places this is found to be nearly filled with native copper, of which the specimen exhibited is one of the masses obtained. An excavation on the vein, 27 feet
DESCRIPTIVE CATALOGUE.

deep, without galleries, yielded about 3 tons of the metal. Other veins intersect the
same rock, and one of these, 6 inches in width, holds good bunches of the variegated sul­
phuret. In ancient shallow holes sunk at intervals along the course of some of the veins
of metallic copper, in this part, there are occasionally found the remains of Indian ham­
mers, consisting of small hammers, usually of trap, having shallow grooves worked around
them, to receive the hammers or thongs attaching the handles, giving evidence of rude
aboriginal attempts at mining, many centuries since.—Quebec group, Lower Silurian.

Smelted.

   a. 2 ingots.

   This copper is smelted at the Bruce mines, from the ore of the neighborhood, and from
   the native copper procured from the 2 feet bed at Michipicoten Island.

Sulphuret.

1. Orford, lot 6, range 12.............................. Geological Survey.
   a. Specimens of the sulphuret of nickel, millerite.

   Associated with a band of serpentine which runs along the east side of Brompton Lake,
in Orford, there is on the lot indicated above, a pale green pyroxenic rock, in which occur
druses, lined with large twin crystals of white pyroxene and with cinnamon colored garnets.
Large masses of calcarespar, probably filling a vein, are here met with, sometimes nearly pure
white and cleavable, at others penetrated and filled with small emerald green crystals of a
chrome garnet. This mineral also forms granular masses, mixed with calcareous spar and
pyroxene, and containing small quantities of the sulphuret of nickel, millerite. Some
specimens of the rock have yielded to analyses as much as 1 per cent. of nickel. The ore
does not appear to be confined to the portion of rock mixed with calcareous spar, but
to penetrate into more homogeneous strata, probably pyroxene, running with the serpen­
tine; where, however, the quantity of the ore seems to be less.—Quebec group, Lower
Silurian.

Native.

   a. Ore from a lode.

   The rock of Prince's location is clay slate, interstratified with greenstone, and overlaid
by a great mass of it with a columnar structure; the whole dipping south-eastward at a small
angle. These rocks are transversely intersected by a vein, which is 20 feet thick on Spar
Island, and from 4 to 5 feet on the main land, running N. W. It is composed of calcarespar,
heavy spar, and amethystine quartz; the latter appearing in druses in the calcarespar. With
these are associated the yellow, variegated and vitreous sulphurets of copper in proming
quantity, with iron pyrites, blende, galena, and silver; the latter occurring both native
and as a sulphuret, in addition to cobalt and arsenic, as well as traces of gold. The
location is the property of the British American Mining Company, and in a small trial
shaft sunk by them to the depth of between 6 and 7 fathoms, on the main land, where the
lode is 4 feet wide, several hundred pounds of the vein, similar to the specimens exhibited
contained 31 per cent. of silver.—Quebec group, Lower Silurian.
MINERALS OF CANADA.

GOLD.

Native.


a. Stream gold in nuggets, 9 among them weighing from 10 dwts. to 126 dwts.

b. Stream gold in dust.

It has long been ascertained that the drift of the south side of the St. Lawrence, in Canada from Lake Champlain to the Etchemin, and probably to the extremity of the province in Gaspé, is auriferous; the area being about 15,000 square miles. Gold has been washed from this gravel on the St. Francis in Melbourne, at Sherbrooke, in Westbury, Woolen and Dudswell, and on Lake St. Francis; as well as on the Chaudière and the Etchemin, and their tributaries, from the sources of these rivers nearly to their mouths. Various companies have made trials of this drift in several places, one of the most important having been on the Rivière des Plantes, in the seigniory of Vaudreuil (Beauce); but of this it is not easy to procure authentic details. In 1851, the Canada Gold Mining Company commenced a trial of the drift along the Rivière du Loup, near its junction with the Chaudière, in the seigniory of Aubert de l'Isle; which continued 3 years. The specimen exhibited is what was obtained by the workings of this Company in 1852, and the following are the results for the years 1851 and 1852:

<table>
<thead>
<tr>
<th>Area</th>
<th>Gold collected</th>
<th>Value</th>
<th>Wages</th>
<th>Profit</th>
</tr>
</thead>
<tbody>
<tr>
<td>4967.30</td>
<td>4223.15</td>
<td>5622.68</td>
<td>690.47</td>
<td></td>
</tr>
</tbody>
</table>

The chief part of the gold was obtained in the bed of the river, but some of it on the bank, and the average thickness of the drift was about 2 feet. The average daily wages were 10 cents a man. The system adopted for dressing was that used in Cornwall for obtaining tin from alluvial deposits.—Drift.


a. Stream gold, a nugget of 80 dwts. with quartz.

In this specimen the proportion of the gold is 64 per cent. It was obtained from the drift of the Rivière des Plantes, a tributary of the Chaudière. Many of the small masses of gold which have been obtained from the drift of the Chaudière valley, being of a character somewhat similar, there cannot be much doubt that the drift gold of the region has been derived from quartz veins, situated probably somewhere not far distant. No quartz so rich in gold as the specimen, has as yet been met with in place in Canada, but the precious metal has been observed in small grains in a quartz vein of between 2 and 3 feet thick, which cuts bluish-black slate, and crosses the Chaudière at the St. Francis rapids, about 3 miles from their foot, and about 3/4 of a mile above St. Francis (Beauce) church.—Drift.


a. Auriferous blonde, galena, arsenical, magnetic and cubic iron pyrites, with quartz and bitter spar; from a vein.

This vein, as just mentioned, is between 2 and 3 feet thick, and consists principally of quartz, in which native gold has been observed; although none is visible to the eye in
the specimens exhibited. The quartz is associated with bitter spar; and in the gangue are disseminated small quantities of galena, blende, arsenical sulphuret of iron, often well crystallized; besides cubic and magnetic iron pyrites. In an analysis made by Mr. Sterry Hunt in 1854, a portion of the galena separated by washing, but still containing a small mixture of the blende and pyrites, gave by assay of 500 grains, 69 per cent. of lead, and 32 ounces of silver to the ton of ore. Another sample of 500 grains, more carefully dressed, gave 37 ounces of silver to the ton. The silver contained a small quantity of gold. Another sample of 500 grains of the sample which gave 69 per cent. of lead, afforded by cupellation, a quantity of silver equal to not less than 226 ounces of silver to the ton. This amount of silver was probably owing to the accidental presence of a fragment of silver ore. 1000 grains of the pyrites, mixed with a little blende, galena, and arsenical pyrites, gave by cupellation 0.16 grains of an alloy of gold and silver. 700 grains of the impure blende gave 0.19 grains of a yellow alloy of the same metals. —Quebec group, Lower Silurian.

   a. Grains in bitter spar.

On the property of Mr. Nutbrown, of which the mining rights have been purchased by the English & Canadian Mining Company, there occurs a vein cutting a bed of steatite. The vein is composed of a gangue of coarsely crystalline bitter spar, mixed with talc, copper glance, and specular iron. There is disseminated, principally in the bitter spar, a small quantity of gold in grains, varying in size from mere points to the magnitude of pin heads. Sometimes the metal appears in laminae in the bitter spar, having a diameter of about 1 or 1/2 of an inch. The vein, which is 2 feet thick, has been mined to a small extent for copper ore. —Quebec group, Lower Silurian.

PLATINUM AND IRIDOSMINE.

Native.

   a. Grains separated from the drift gold.

Among the drift gold of the Chaudière there are met with, in very small quantities, grains of platinum, and of iridosmine, the latter being an alloy of the rare metals iridium and osmium, which being very hard, is used for pointing gold pens. Some of the gold met with on the Chaudière has been found thinly coated with a mercurial amalgam; but no trace of cinnabar, the commonest form of the ore of mercury, has been observed in the drift. Among the substances met with by the Canada Gold Mining Company, in separating the gold from the drift, lead shot of various sizes, from partridge to swan shot, were nearly as abundant as the gold. —Drift.
MINERALS APPLICABLE TO CHEMICAL MANUFACTURES.

Chromic Iron.

   a. Specimens from the surface.
   Loose masses, traced for half a mile, running with the strike of the serpentine, of which
   the mountain is composed; the largest masses weighing about 20 pounds.—Quebec group,
   Lower Silurian.

   a. Specimen from a bed.
   A bed of about 14 inches thick in serpentine. The bed has been partially worked by
   the proprietor, Mr. Leckie, of Acton Vale; who obtained about 10 tons of ore, with 46 per
   cent. of oxyd of chromium, from 7 square fathoms in the plane of the bed. The ore occu-
   pied a lenticular form in the serpentine.—Quebec group, Lower Silurian.

   a. Specimen from a bed.
   A bed of from 12 to 24 inches thick in serpentine, on the property of Mr. L. A. Perkins.
   The bed dips to the eastward, at an angle of about 80°, and the ore occurs in masses of
   from 50 to 1000 pounds in weight.—Quebec group, Lower Silurian.

   a. Specimen from a bed.
   A bed of uncertain thickness in serpentine, in which the ore runs in lenticular nodules
   of from 6 to 9 inches thick.—Quebec group, Lower Silurian.

Molybdenite.

1. Quetachoo River, Manicouagan Bay, N. shore Gulf
   of St. Lawrence. Geological Survey.
   a. Specimens from a bed.
   A bed of 6 inches thick in gneiss, in which the sulphuret of molybdenum occurs in
   nodules of from 1 to 3 inches in diameter, and in flakes of ¹₄ or ⅓ of an inch thick, and
   12 inches in diameter.—Laurentian.
DESCRIPTIVE CATALOGUE.

Cobaltiferous pyrites.

1. Elizabethtown, near Brockville. \textit{A. S. Brown, Brockville.}
   \textit{a. Specimens from a bed.}
   A great irregular mass in gneiss, probably lenticular, running with the stratification.
   It has been excavated to a breadth of 20 feet, but its length and full thickness have not
   been ascertained. \textit{Assays of the ore have yielded} 1 per cent. of cobalt. This according to
   the newest methods of extraction of McFarlane, Roscher and Dahl, would yield a pro-
   fitable result. \textit{The ore is on the property of Mr. Billings.—Laurentian.}

Dolomite.

1. Brome, lot 16, range 11. \textit{Geological Survey.}
   \textit{a. Specimens from a bed.}
   In the Eastern Townships a vast quantity of dolomite occurs. Stratigraphically, it is at the
   base of the Quebec group, where magnesian rocks of different descriptions are associated
   with the sulphures of copper and other metals. The dolomite occurs in bands, which are
   from 100 to 300 feet thick. These run parallel to one another, on the opposite sides of
   synclinal and anticlinal forms, by which the bands are repeated in many places. The
   exposure in Brome, from which the specimen exhibited is taken, is on the east side of the
   Shipton and St. Armand anticlinal, and has been traced for many miles, running N. E.
   and S. W.—Quebec group, Lower Silurian.

Magnesite.

1. Sutton, lot 12, range 7. \textit{Geological Survey.}
   \textit{a. Specimens from a bed.}

   \textit{a. Specimen from a bed.}
   One of the rocks associated with or replacing the dolomite, at the base of the Quebec
   group is magnesite. It is of more rare occurrence than the dolomite. In Sutton, it occurs
   on the east side of the Shipton and St. Armand anticlinal, where it is often slaty, from a
   mixture of feldspar and a mica colored green by chromium. The purest specimens
   contain 80 per cent. of carbonate of magnesia, with a portion of carbonate of iron. In
   Bolton, it occurs on the east side of the Melbourne and Potton anticlinal, where it forms
   an enormous bed, resembling a crystalline limestone. It contains, like the last, small por-
   tion of chromium and nickel, and consists of:
   \begin{align*}
   \text{Carbonate of magnesia,} & \quad 60 \\
   \text{Carbonate of iron} & \quad 9 \\
   \text{Grains of quartz} & \quad 31 \\
   \hline
   \text{100}
   \end{align*}
   Though the use of this mineral as an economic source of pure magnesia and magnesian
   salts on a large scale, may be worthy of consideration, its most important application is
   probably for the fabrication of a cement to resist the action of sea-water.—Quebec
   group, Lower Silurian.
MINERALS OF CANADA.

Petroleum or rock oil.

1. Enniskillen, lot 14, range 2. [Canadian Oil Company, Hamilton.]
   a. Gum or mineral tar from the surface.
   b. Crude oil from the well.
   c. Refined or burning oil.
   d. Lubricating or machinery oil.

Natural springs of rock oil have long been known in several localities in Western Canada. Two of these are in the township of Enniskillen, in the southern part of which are two patches of an acre or more, covered with a layer of several inches of viscid mineral tar or asphaltum, which has resulted from the drying up of the petroleum of these springs. Wells sunk in their vicinity, to a depth of from forty to sixty feet, through the superficial clays, encounter a stratum of gravel, resting on the surface of the rock beneath, and often filled with oil; giving origin to what are called surface wells. On boring into the underlying soft fossiliferous shales and limestone, fissures are met with at various depths, from which rise abundant supplies of oil, often accompanied with inflammable gas, and with water, which is sometimes saline. These fissures, which also supply the surface wells, are apparently connected with the oil-bearing strata of the Corniferous limestone; which is from 200 to 300 feet below the surface, in Enniskillen. Within an area of about four square miles in the first three ranges of this township, there were supposed to be, in August, 1861, about seventy wells, yielding more or less oil. Of these, 40 were surface wells, that is, wells sunk from 40 to 60 feet, through the drift clay and gravel, to the rock beneath. Some of these latter, which had yielded but little oil, gave abundant supplies by boring into the rock. The oil-bearing fissures or veins, in adjacent wells, were met with at depths varying from 36 to 100 and even 150 feet from the surface of the rock. One of the most abundant occurred at 60 feet. In some few cases the oil from the borings rises above the surface of the ground, constituting what are called flowing wells.

It is not easy to know the amount of oil which these wells are capable of supplying; since from the great difficulty in getting it to market, arising from the want of good roads, few of the wells are regularly and continuously pumped. Some of those which were bored in July and in August last, are said upon good authority, to have yielded from 400 to 600 barrels of oil, in a week or two after having been opened; but the reservoirs provided, being filled with oil, the pumping of the wells was suspended. Two bored wells, belonging to Mr. Williams, which were the only ones continuously wrought in August last, are said to have yielded together, during some months, from 20 to 25 barrels (of 41 gallons each) daily. About 6 miles to the northward, at Petrolia, on lots 13 and 14 of ranges 10 and 11 of the same township, 16 wells had been sunk last August; of which 12 were surface wells, and had yielded large quantities of oil. Several of these had been wrought for nearly twelve months, and were supposed in that time to have yielded 1000 barrels. Other wells had recently been bored to a depth of nearly 200 feet, but yielded less oil than the surface wells. The wells of this region seem, thus far, to be less important than those in the southern part of the township. The oil from the deep or rock is somewhat lighter and more fluid than that from the superficial wells, which is very dark colored and somewhat viscid. Great expectations have recently been excited by a flowing well, known as Shaw's, which was sunk to a depth of about 200 feet, and when first opened, a few weeks since, was supposed to yield, for a short time, 2000 barrels of oil in 24 hours; which flowed into a stream near by and was lost. This well is however said to have been since closed, so that the discharge is under control. Another recent well, near by, known as Bradley's, is nearly as abundant. The experience in Pennsylvania has however shown that the supply from these flowing wells soon diminishes, and eventually fails. Adjacent borings sometimes appear to be connected with the same oil-yielding fissure, and to affect each other's supply; in some cases air passes down one shaft when the other is pumped. — Corniferous formation, Devonian.
2. Tillsonburgh............................ Watkins and Inglis, Hamilton.

a. Crude oil from the well.

Near the village of Tillsonburgh, in the township of Dereham, natural oil springs occur, and two wells have been bored in the Devonian limestone, which is here covered by about 40 feet of clay and sand. One of these had been sunk 36 feet in the limestone, and had furnished when seen, in August, a few barrels of oil.

In the townships of Zone, Moss, and Orford, on the banks of the Thames, oil springs abound for a distance of about 4 miles. These, like the other natural springs mentioned above, furnish but small quantities of oil; several wells have however been sunk in the clay, and the rock beneath has been drilled. One of these, at a depth of 70 feet in the clay, had yielded about 40 barrels of oil.—Corniferous formation, Devonian.


In a quarry on the lot mentioned, 2 oil-bearing beds, one of 2 and another of 6 inches, are seen: they are made up in great part of corals of the genera Heliothyllum and Faversites, in the pores of which the oil is lodged like honey in a comb. Other coral beds in the same series, however, are quite free from oil. The limestone beds above and below these are compact, and not at all impregnated with oil, which, even in the coral beds, is seen, when these are freshly broken, to be confined to the fossils, and not to be uniformly disseminated in the layer. When the rock is quarried, the oil flows out, and may be collected on the water in the bottom of the quarry. The facts observed with regard to the petroleum springs in Canada and the United States, would seem to show that they are always on the lines of anticlines, along which the oil from its superior levity accumulates, and afterwards, by the pressure of water, is raised to the surface through the natural fissures which generally occur upon anticlines. The oil-bearing limestone underlies an area of 7000 square miles in Western Canada. This limestone is of marine origin, and contains no organic remains but those of marine animals; so that we are led to conclude that these hydrocarbons have been derived from a peculiar decomposition of their tissues. These tissues, as is well known, differ but little from those of the plants, which in many more recent formations have given rise to bitumens. We may suppose that many soft gelatinous animals, and perhaps plants, whose traces have disappeared, may have contributed to form the petroleum of these coral beds.—Corniferous formation, Devonian.

Bituminous shale.


a. Shale from the bed.
b. Burning oil.
c. Lubricating oil.

The shale of Collingwood, on lot 23, range 3, yields, when distilled, from 3 to 4 per cent. of tarry oil, which by the usual process of rectification affords oil fit for illumination and lubrication. Works were erected by McDermot, Pollard & Macdonell, in October, 1859, consisting of 24 retorts, and capable of yielding about 250 gallons of oil daily, by the distillation of from 20 to 30 tons of shale. The available bed of shale is 7 feet in thickness, and the material was delivered broken for the retorts, at 20 cents the ton. The cost of the crude oil was said to be 14 cents the gallon, and for a while the works were carried on successfully, a ready market being found for the oils; but the works were repeatedly destroyed by fire, and the oils from this source coming in competition with petroleum from the oil wells of Enniskillen, the enterprise is for the present abandoned.—Utica formation, Lona Sillurian.
Phosphate of lime (apatite).

   a. Specimen from a bed.

   This deposit has been traced across lots 24, 25, and 26, range 8 of North Elmsley, for a
distance of about a mile, in a direction nearly S. W. It apparently forms an irregular bed
in the Laurentian limestone. On lot 25, where it has been somewhat quarried, the
breadth of the bed seems to be about 10 feet, of which 3 feet are nearly pure crystal-
line apatite, with only a small admixture of black utica. The remainder is mingled
with the limestone, the phosphate, however, in many parts largely predominating.—Laure-
rentian.

2. South Burgess, lot 9, range 5. ...................... F. Poole, Perth.
   a. Specimen from a bed.

   The deposit of phosphate of lime in Elmsley appears to be continued south-westwardly
through Burgess. Indications of it occur on lot 2, range 7, and the quantity on lots 7, 8,
9, 10, range 5, still farther on, appears to be important.—Laurentian.

3.

REFRACTORY MINERALS.

Soapstone (steatite, compact talc).

   a. Cut specimens.

   a. Cut specimens.

   Among the magnesian rocks at the base of the Quebec group, in that part of its dis-
tribution where it is in a metamorphic state, soapstone or steatite occurs in great abun-
dance. Beds of it, varying in thickness from 1 to 15 feet, can be traced for long dis-
tances, usually not far removed from serpentine, dolomite, or magnesite; or apparently
replacing one or other of these rocks. In general the soapstone is remarkably pure, but
occasionally there are disseminated in it crystals of bitter spar or of actinolite. The spe-
cimens exhibited from Sutton and Bolton are from equivalent bands of 20 and 30 feet
respectively, on the opposite sides of Sutton Mountain. In the latter locality the soap-
stone is interstratified with potstone and dolomite, and in some parts of the band, the three
rocks are seen to interlock among one another in lenticular masses. These two bands of soap-
stone appear to be on the opposite sides of a general synclinal form; yet Sutton Mountain
between them, has an anticlinal structure, with a height stated to be four thousand feet.
This mountain occupies a breadth of 10 miles at the province line, but gradually narrows, and completely dies down in a distance of 80 miles north-eastward. Its structure may be explained by stating that Sutton valley on the west, Sutton Mountain in the middle, and Potton valley on the east, run upon three anticlinal axes, which converge to the north-eastward, like the sticks of a fan, and while the rocks on the two outside anticlinals have been worn into valleys, those on the middle anticlinal have resisted denudation. Sutton Mountain is continued into Vermont, in Jay Mountain; which appears to stand on one of the main axes of the Green Mountain range.—Quebec group, Lower Silurian.

**Potstone (compact chlorite).**

   a. Cut specimens.

   A considerable portion of the rocks of the Quebec group, in their metamorphic condition, consists of chloritic slates; which appear to occupy a somewhat higher stratigraphical place than the more magnesian strata just mentioned, and usually to fill up the middle, and more elevated parts of the synclinal forms, in which the Quebec series is distributed through the country. There occur also bands of pure compact chlorite or potstone interstratified with the more magnesian strata. Some of these are of considerable thickness, and the one in Bolton, from which the specimens are taken, has a width of about 20 feet.—Quebec group, Lower Silurian.

**Mica rock.**

   a. Uncut specimens.

   In nearly the same stratigraphical place as the potstone, there occurs in some localities a compact, hydrous mica, which so much resembles potstone as to have been mistaken for it: and very probably it possesses the same refractory properties. Where the specimens were obtained, a breadth of 5 feet is exposed; the full thickness of the band, however, is supposed to be much greater.—Quebec group, Lower Silurian.

**Mica.**

   a. Cut and dressed plates.

   a. Specimen in crystals.

   a. Cut and dressed plates.

4. South Burgess, lot 1, range 4 .......................... A. Cowan, Kingston.
   a. Uncut specimens.
Magnesian mica or phlogopite occurs abundantly, in small scales, in the crystalline limestones of the Laurentian system, but sometimes also in crystals sufficiently large to be economically available. These are generally met with near bands of quartzite or of pyroxenic gneiss limiting the limestones or near some interstratified mass of a similar character, and they are usually associated with other minerals. Among these, in addition to quartz, pyroxene and feldspar, there occasionally occur tabular spar, apatite, sphene, iron pyrites, idocrase, garnet, tourmaline, zircon, and sometimes corundum. In Grenville, where the mineral is imbedded in massive pyroxene rock, close alongside of a band of crystalline limestone, crystals of mica have been obtained, giving sheets measuring 24 by 14 inches. In North Burgess, where it has been mined by Mr. Cowan, on lot 17, range 9, the mica is imbedded in a soft pyroxenic rock, running apparently N. E. and S. W., and limited by a band of quartzite on the southward side. The mica here appears to run for 75 yards in pretty regular bands, and some of the sheets, after being dressed, are as much as 20 inches square; some have been obtained measuring 20 by 30 inches.—Laurentian.

Plumbago.


   a. Specimen from a bed.


   a. Specimen from a bed.


   a. Specimen from a bed.

The crystalline limestones of the Laurentian system are marked almost universally by the occurrence of graphite or plumbago, in small scales, which are often so thickly disseminated in particular bands of the rock, as to give them a black or dark grey color, distinctly marking its stratified character. It occurs also in beds, in sufficient purity and quantity to be economically available. The workable beds which have been observed, are chiefly in various townships on the north side of the Ottawa. They occur in many localities, at considerable distances from one another, but several of the exposures are probably repetitions of the same bed, or, at any rate, of beds approximately equivalent, in repetitions of the same band of limestone. The whole Laurentian series is so corrugated, that the outcrop of one of these bands of limestone, in the counties of Argenteuil and Two Mountains, followed through all its windings, in an area of 50 miles northward by 20 miles eastward, measures upwards of 200 miles. A bed of pure graphite occurs in the Augmentation of Grenville township, and has been traced at intervals, for a distance of about three miles, running a little east of north. One of the exposures, occurring on lot 3, range 2, has, to a small extent, been mined by Messrs. Russell & Co. At the opening of the excavation, it showed a thickness of about 3 inches, but the pure graphite was found to form a lenticular mass, which appeared to be separated from other masses of the same character by intervals, in which the graphite became intermixed with the limestone. It is probable that a number of these, running through the rock at the same horizon, may represent the general character of the workable beds. On lot 3, range 6, the bed becomes 3 feet thick, but here the quality is impaired by the presence of foreign earthy matters, which, however, can scarcely be detected by the eye.—Laurentian.
Asbestus.

   a. Specimen from a vein.

   Asbestus, generally a fibrous serpentine or chrysotile, occurs in veins cutting the serpentine of the Eastern Townships.—Quebec group, Lower Silurian.

Friable Sandstone.

1. Pittsburgh, lot 20, range 1, ......................... A. Cowan, Kingston.
   a. Specimen as sent to foundries.

   This crumbling sandstone occurs in a bed upwards of 20 feet thick, and is in much demand for iron foundries, being used to protect the sides and bottoms of the furnaces. It is supplied to the founders of Montreal at $3, and those of Toronto at $2.50 per ton after being carried about 170 miles in opposite directions to each place. About 1500 tons are consumed at the foundries of these two cities.—Potsdam formation, Lower Silurian.

Fire-Clay.

1. Dundas ................................................. Geological Survey.
   a. Specimen of the clay.

   This clay is derived from an argillaceous band 20 feet in thickness, near the base of the Clinton formation. The rain washes the clay from the bank, and deposits it in the bottom of pools at its foot. When the water dries up in these, the clay is dug from them, and is used in the iron foundries at Dundas and at Hamilton. From the neighborhood of Dundas, the formation strikes south-westward, and after folding over an anticlinal axis, which runs southward of west from the extremity of Lake Ontario, it returns on the south side of Lake Ontario, towards the Niagara River. The same clay band is thus again met with in the escarpment near the village of Ancaster. The clay has only lately come into use, and not much is yet known regarding its merits, but it is said to answer a good purpose, and in Mr. Gartshore's foundry, at Dundas, has entirely superseded the fire-clay formerly imported from the United States.—Clinton formation, Middle Silurian.

2. St. Foy, near Quebec, ......................... Michael Finley, Quebec.
   a. Specimen of clay.
   b. Piece of pottery made from the clay.

   This clay, which is represented by the contributor to be of a refractory character, forms a considerable deposit at Belmont, the property of Mr. J. W. Dunscomb.—Drift.
MINERALS APPLICABLE TO COMMON AND DECORATIVE CONSTRUCTION.

BUILDING STONES.

Limestones.

1. Arnprior, MacNab, lot 4, range C, Geological Survey.
   a. 1 foot cube dressed.
   This building stone is derived from one of the bands of crystalline limestone of the Laurentian series. It occurs on the property of Mr. McLaren, close upon the margin of the Lac des Chats, one of the expansions of the Ottawa, where a great supply of the rock might be obtained. It has been used by the Board of Works in the construction of a bridge over the river Madawaska, the mouth of which is near the limestone.—Laurentian.

2. Phillipsburg, St. Armand, C. R. Chezzeman, Phillipsburg.
   a. 1 foot cube dressed.
   The exposure of limestone from which this stone is derived, occurs within a mile of Phillipsburg, Missisquoi Bay, on the land of the exhibitor, near the line of a proposed railway. The rock is compact and crystalline, dresses easily, and appears to have considerable strength. A few obscure fossils are met with in the rock, belonging to the genera Pleurotomaria and Holopea. Higher in the series, the organic remains are more distinct, and shew the formation to be equivalent to the Calciferous.—Quebec group, Lower Silurian.

   a. 1 foot cube dressed.

   a. 1 foot cube dressed.

   a. 1 foot cube dressed.
   The specimens of building stone from Caughnawaga, St. Dominique, and East Hawkesbury, 3-5, are all from one geological formation, the Chazy; which in the area indicated by the distribution of the places named, is composed of massive beds, yielding blocks of stone fitted for the purposes of canal locks and railroad bridges. The quarries of Caughnawaga have supplied a large amount of material for the upper locks of the Lachine Canal, and those of the Beauharnois Canal. That of Hawkesbury, as well as a quarry on the same formation on Isle Bizard, has furnished blocks for the Carillon Canal. The same formation, near Montreal, and on Isle Jesus, near Terrebonne, has been resorted to for similar blocks for the lower locks of the Lachine Canal. In all the places named, the beds abound
in the remains of encrinites and cystidians, and the peculiar crystallization of these, give
to the rock a crystalline texture, which constitutes one of its valuable characters. There is
occasionally some admixture of magnesia in the stone, and after long exposure to the
atmosphere it assumes a yellowish tinge. The St. Dominique stone is more compact than
that of Caughnawaga, and has been used for the purposes of the St. Lawrence and Atlantic
Railway.—Chazy formation, Lower Silurian.

a. 1 foot cube dressed.

a. 1 foot cube dressed.
The stone of Pointe Claire and Cornwall, 6–7, belongs to the formation next above the
Chazy. It is a black compact and thick bedded. At Mille Roches, near Cornwall, it
has supplied stone for the construction of the locks of the Cornwall Canal, and at Pointe
Claire, it has furnished the stone used in building the piers of the western half of Victoria
Bridge, while that used for the other half was brought from what is said to be the same
formation, at Île la Motte in Lake Champlain. On the north side of the Pointe Claire
quarry, there is a vertical exposure of 30 feet of massive strata, varying in thickness from
1 to 3 feet, the blocks obtained from which, for the bridge, ranged from 4 to 7 tons.—Birds·
eye and Black River formation, Lower Silurian.

a. 1 foot cube dressed.

a. 1 foot cube dressed.
The Trenton formation, which is the next in succession above the Birdseye and Black
River, yields excellent building stone at Montreal, at Chevrotière, nearly 40 miles above
Quebec, and at many intermediate places. The best stone at Montreal is derived from a
10 feet band of grey bituminous granular limestone, in beds of from 3 to 18 inches thick at
the bottom, passing at the top, into a black nodular bituminous limestone: which is inter·
stratified with black bituminous shale, in irregular layers of from 1 to 3 inches. This grey
limestone, which is near the base of the formation, is a mass of comminuted organic
remains, which consist largely of the ruins of crinoids and cystidians. The crystallization
of these fossils gives a crystalline character to the rock. A considerable number of quar·
ries are worked upon this band of grey limestone, there being 4 principal ones near Mon·
treal, and the best houses of the city are built of the stone. The quantity of stone annually
quarried in the immediate vicinity of Montreal is computed to be:

<table>
<thead>
<tr>
<th>Cubic Feet</th>
<th>Tons</th>
</tr>
</thead>
<tbody>
<tr>
<td>313,200</td>
<td>28,600</td>
</tr>
</tbody>
</table>

The prices of good stone in Montreal are:

<table>
<thead>
<tr>
<th>Stone Type</th>
<th>Price</th>
</tr>
</thead>
<tbody>
<tr>
<td>Ashler stone, undressed</td>
<td>$0.18 per square foot</td>
</tr>
<tr>
<td>Ashler stone, dressed</td>
<td>$0.30</td>
</tr>
<tr>
<td>Mouldings, from $0.16 to $3.00, or for a fair moulding, $1.50 per linear foot.</td>
<td></td>
</tr>
<tr>
<td>Fluted columns 15 inches diameter, for the stone, $1.00 per rising foot.</td>
<td></td>
</tr>
<tr>
<td>Heavy rough stone, from 6 to 30 cubic feet, from</td>
<td>$0.30 to $0.60 per cubic foot.</td>
</tr>
<tr>
<td>Very heavy rough stone, say 60 cubic feet,</td>
<td>$1.00 per cubic foot.</td>
</tr>
</tbody>
</table>
The strata in the neighborhood of the city are much traversed by trap dykes, which probably have a connection with an intrusive mass extending over 700 acres, and constituting Mount Royal, from which the city and island take their name. Some of the quarries display a number of these trap dykes, which run in several directions and intersect one another. In some instances, the limestone, having been removed from among them, the dykes are left standing up several feet above the bottom of the quarries, representing in a marked manner the various details of the cracks they once filled.

In the seigniory of La Chevrotière, a very excellent limestone for building is obtained between 3 and 4 miles back from the St. Lawrence. It usually goes, however, under the name of the Deschambault stone, in consequence of it being put on board of boats at this place. The stone is of a yellower or warmer grey than the Montreal stone; it is more even in its tint, and becomes somewhat less discolored by weathering. It is more granular and more easily cut, being softer and tougher, but it does not take so fine nor so sharp an edge, nor does it pick so well. Three beds of pretty uniform character are worked; the top and bottom ones are 18 inches thick each, and the middle one 3 feet. There is said to be a fourth bed beneath, with a thickness of 4 feet, which has not been quarried. The strata are so nearly horizontal, that it is difficult to determine their dip; it is therefore probable that the stone will spread to a considerable extent in the vicinity. Along the concession line it is known for 26 acres to the S. W., and 5 acres to the N. E., and on the road across the concession it is visible for a breadth of 10 acres; beyond which, in sinking wells to a depth of 20 feet in blue clay, no rock is met with. The produce of the quarries of La Chevrotière has a deserved celebrity in Quebec, where it has been used in the construction of churches and other buildings.—Trenton formation, Lower Silurian.

Dolomites.


   a. 1 foot cube dressed.

   This beautiful and enduring stone can be obtained in unlimited quantities, the formation from which it is derived being here 150 feet in thickness and divided into beds varying from a few inches to 9 feet. This stone possesses the very great advantage of being free from any substance producing stains. Its color rather improves with the weather, and the beauty of no building erected of it appears as yet to be marred by the growth of lichens. It is specially adapted for heavy masonry, and blocks of any required size can be obtained. The quarries are about 1/4 a mile from the harbor.—Niagara formation, Middle Silurian.


   a. 1 foot cube dressed.

   This stone is from the lower part of the Niagara formation, and is rather more compact than the Owen Sound specimen. The cliff is here about 50 feet high, and might be quarried with the greatest facility. Few of the beds are less than 2 feet in thickness, and some of them are about 5 feet, but the locality is not near any navigable water or railway.—Niagara formation, Middle Silurian.
3. Rockwood, Eramosa, lot 5, range 4 ...................... Geological Survey

   a. 1 foot cube dressed.

   This specimen is also from the Niagara formation, which is here more than 100 feet thick. The greater part of it consists of thick bedded light grey porus crystalline dolomite. The beds vary from a few inches to 10 feet in thickness; about 30 feet of it is almost white. Buildings of cut stone obtained from this band, are observed to improve in color after exposure, and have a silvery white appearance at a short distance. The piers of the long railway viaduct over the valley of the Eramosa, at Rockwood, are built of stone from this formation, and have a very substantial appearance. The axis of an east and west anticlinal form runs under Rockwood, carrying a spur of the Niagara formation several miles to the eastward of the general trend of the outcrop. A north and south anticlinal passes under the same place, being one of a series which produces southward indentations in the outcrops of the palaeozoic strata all the way from Kingston to the main body of Lake Huron.

   —Niagara formation, Middle Silurian.


   a. 1 foot cube dressed.

   This stone is from the immediate vicinity of the thriving town of Guelph. The quarries expose 15 feet of strata similar to the specimen. The thickest bed is 4 feet, and the thinnest about 3 inches. The stone is a light grey crystalline dolomite, like the last, somewhat cellular, but strongly coherent. It is easily worked, is suitable for the best architectural purposes, and appears to be of a durable character. The Guelph formation extends over a large area, and much of it is of the same character as the specimen.—Guelph formation, Middle Silurian.

5. Oxbow, Saugeen River, Brant, lot 2, range 8......... Geological Survey.

   a. 1 foot cube dressed.

   b. Journal bearer from a lower bed.

   The beds from which the block a is obtained, are near the summit of the Onondaga formation, and yield probably the best dolomite for fine architectural purposes which has yet been discovered in the country. It resembles the Caen stone in the facility with which it can be worked, but it is closer grained, and by no means so absorbent, and is thus better adapted for withstanding the Canadian climate. Two bands, of about 10 feet each, occur here in the upper part of the Onondaga formation. The higher one is exposed at the surface, in a position offering every facility for quarrying it. The bed from which the specimen was procured, is 2 feet thick, free from stains, and splits with great precision with the plug and feather. The whole upper band is composed of thick beds of the same character; the thickest one in the lower band measures over 3 feet. The locality is near a projected line of railway, and is 22 miles from Southampton harbor by the present road. It overlooks the Saugeen River, down which large scows can be floated to Southampton.

   The specimen b is from a very light grey oolitic bed, 17 inches thick, immediately beneath the previous bed; it has been used for supporting water wheels, in mills in the neighborhood, and found to answer well, becoming highly polished under the action of a revolving shaft.—Onondaga formation, Upper Silurian.
Sandstones.

1. Lyn, Elizabethtown, lot 26, range 2 ........................... Geological Survey.
   a. 1 foot cube dressed.

   a. 1 foot cube dressed.

   a. 1 foot cube dressed.

   a. 1 foot cube dressed.

   These specimens, 1-4, are derived from the Potsdam formation, constituting the summit of the lowest group of fossiliferous rocks of Canada. A quarry has been opened on the outcrop of the rock, at Lyn, by Mr. B. C. Brown, and the stone from this and from a quarry on the property of Mr. Keefer, Nepean, in the same formation, has been used in the construction of the new Parliament buildings at Ottawa. At Lyn, the beds of the formation are massive, and are seen resting on the Laurentian gneiss.—Potsdam group, Lower Silurian.

5. Pembroke...................................................... Geological Survey.
   a. 1 foot cube dressed.

   This fine freestone is much exposed in the vicinity of the Allumette rapids, near Pembroke. A quarry has been opened on it, on the land of Mr. Peter White, where it occurs in beds varying in thickness from 6 to 18 inches. It is easily wrought and carved, and though soft, it is tough, and retains sharp angles and corners.—Chazy formation, Lower Silurian.

   a. 1 foot cube dressed.

   This fine grained compact greenish-gray sandstone is from a deposit of about 10 feet in thickness. Some of the beds are thick, but some thin enough for flagstones; the stone is free from iron stains, but subject to a growth of lichens in shaded and moist situations. —Gray band, Medina formation, Middle Silurian.

   a. 1 foot cube dressed.

   This is from a bed of light grey freestone, occurring in a band of about 20 feet in thickness. The beds are mostly thick, fine grained and compact; some split into good flagstones; but all are rather hard for grindstones. It has been used in constructing culverts on the Grand Trunk Railway and numerous buildings in Toronto; among them the University and other important structures, and appears to answer well.—Gray band, Medina formation, Middle Silurian.
   a. 1 foot cube dressed.
   b. 1 foot square, pierced for stove pipe.
   These specimens are from a band of fine grained soft light grey freestone, supposed to be
   20 feet thick. The beds are from 2 inches to 3 feet in thickness; some of them reedy, or
   marked by lines of stratification. The stone yields good grindstones, but has not yet been
   much used for building purposes, although from the specimen a, it would appear to be well
   suited for such. From the facility with which parallel faced blocks of the required thick-
   ness can be obtained, this stone is well adapted for stove-pipe holes, for which it is much
   used.—Grey band, Medina formation, Middle Silurian.

   a. 1 foot cube dressed.
   A band of white sandstone runs through Haldimand County in Western Canada, and
   is largely developed on the Oneida and North Cayuga town-line, north of the Talbot road.
   Its beds are massive, ranging in thickness from 1 to 3 feet, and when fine grained, it is well
   adapted for building purposes. A quarry has been opened in it on the land of Mr. Wil-
   liam De Cew, from whom this specimen of building stone was obtained.—Oriskany for-
   mation, Devonian.

Labradorite.
   a. 1 foot cube dressed.
   The opalescent variety of labradorite occurs in cleavable masses in a fine grained base of
   the same mineral character, composing mountain masses. Where these are thickly dissemi-
   nated in the paste, the stone would become a beautiful decorative material, applicable to
   architectural embellishment, and articles of furniture. Its hardness is about that of ordinary
   feldspar, and it would, in consequence, be more expensive to cut and polish than serpen-
   tine or marble, but it is not so readily scratched or broken, and would therefore prove much
   more lasting. Professor Emmons states that a block of the stone submitted to the action of
   a common saw, such as is used in sawing marble, moved by the waste power of a com-
   mon water mill, was cut to the depth of 2 inches in a day: which is understood to be one-
   fifth the amount that would be cut in a block of good marble, in the same time, by the
   same means. It would thus appear that though the operation is slower in the case of
   labradorite, there is no greater amount of mechanical contrivance required than for mar-
   ble, and that slabs could be prepared for chimney pieces, for pier tables, and other articles
   of furniture, at a cost beyond that of marble, not greater than is proportionate to the supe-
  rior beauty and durability of the material.—Laurentian.

Gneiss.
   a. 1 foot cube dressed.
   This stone has been used for building the dam and reservoir of the Quebec water-works,
   on the St. Charles river. The gneiss, which is obtained a short distance above the reser-
   voir, is hornblendic, being composed of translucent, colorless quartz, white orthoclase, (the
feldspar predominating over the quartz) and black hornblende, all running in irregular parallel planes, showing the gneissoid structure very distinctly, and having at a little distance, a general grey color. The rock splits in almost any direction by means of wedges, but most easily in that of the gneissoid layers, particularly when these are even. The layers are, however, occasionally affected by undulations and contortions, but these do not materially affect its dividing by means of wedges. The rock splits and dresses with most difficulty at right angles to the gneissoid layers. It is capable of receiving fine, smooth faces, giving sharp edges and corners. Masses of almost any size can be blasted out from the rock, and large blocks have been dressed and applied to the masonry work of the reservoir, which will no doubt prove a structure of the most lasting character.—Laurentian.

a. 1 foot cube dressed.

The porphyroid orthoclase gneiss, which this specimen represents, forms great mountain ranges among the Laurentian rocks, rising into the highest peaks of the orthoclase region, and generally constitutes the main body of rock, separating one important band of limestone from another. These masses appear sometimes to attain several thousand feet in thickness, divided however at unequal intervals, by thinner and less feldspathic bands, in which the stratification is more distinct.—Laurentian.

Syenite.

a. 1 foot cube dressed.

a. 1 foot cube dressed.

3. Barrow Island, River St. Lawrence, opposite Gananoque .... Geological Survey.
a. 1 foot cube dressed.

The intrusive masses of the Laurentian series consist chiefly of syenite and dolerite. They occur in many parts of the country, but their relative ages have been ascertained almost altogether by the investigation of the counties of Ottawa and Argenteuil. What appear to be the oldest intrusive masses are a set of dykes of a rather fine grained dark greenish-grey greenstone or dolerite, varying in thickness from a few feet to 100 yards. Their general bearing appears to be E. and W. These greenstone dykes are interrupted by an intrusive syenite, a mass of which occupies an area of about 36 square miles in the townships of Grenville, Chatham, and Wentworth. The specimens 1, 2, are derived from it, and 3 is from a mass of a similar character, occurring between Kingston and Gananoque. In Grenville the syenite is cut and penetrated by dykes of a porphyritic character. These masses belong to what has been called felsite porphyry, hornstone porphyry, or orthophytre, having for its base an intimate mixture of orthoclase and quartz, colored by oxyd of iron, and varying in color from green to various shades of black. Throughout the paste, which homogeneous and conchoidal in its fracture, are disseminated well defined crystals of
a rose-red or flesh-red feldspar, apparently orthoclase, and, although less frequently, small grains of nearly colorless quartz. All of these intrusive masses are cut by another set of dolerite dykes, which probably belong to the Silurian period, or perhaps to the Devonian.

—Laurentian.

Granite.

   a. 1 foot cube dressed.

   This band of granite, which has a considerable proportion of quartz, has been used in the seigniory of St. Joseph for millstones, and would yield a strong and durable building stone, is about 50 or 60 feet thick. It runs with the stratification, near to a band of serpentine, and is supposed to be an altered and not an intrusive rock.—Quebec group, Lower Silurian.

   a. 1 foot cube dressed.

   An intrusive granite of Devonian age occurs in considerable abundance in the Eastern Townships, and forms many isolated hills, the whole of them of small size, with the exception of Great Megantic Mountain, which occupies an area of about 12 square miles. The rock splits well with plug and feather, and can be obtained in blocks of almost any required size. It forms a handsome building stone, and has been used for bridges on the St. Lawrence and Atlantic Railway. It is composed of white quartz and white orthoclase feldspar with black mica. An area of this rock occurs in Stanstead, covering 6 square miles, and forming Riby plains. Another occurs in Barnston, from which the specimen now exhibited was obtained. Granite of the same character, and probably of the same age is widely distributed in the State of Maine, and is traceable to New Brunswick, where it is overlaid by the Carboniferous rocks.—Devonian.

MARBLES.

Limestones.

   a. Striped light and dark grey, large pattern.
   b. "", "small pattern.
   c. "", "cut across the beds.

   At the mouth of the Madawaska, in McNab, an almost unlimited amount of crystalline limestone is marked by grey bands, sometimes narrower, and sometimes wider, running in the direction of the original bedding, and producing a regularly barred or striped pattern, where there are no corrugations in the layers. When the beds are wrinkled, there results a pattern something like that of curly grained wood, when smooth planed. The colors are various shades of dark and light grey, intermingled with white. The shades arise from a greater or less amount of graphite, which is intimately mixed with the limestone. The granular texture of the stone is somewhat coarse, but it takes a good polish, and gives a pleasing marble. Mr. W. Knowles has opened a quarry in limestone of this description at Arnprior, and erected a mill for the purpose of sawing and polishing it, for chimney pieces, monuments, and other objects. A monument of it has been erected in the Mount Royal cemetery.—Laurentian.
2. Elzevir ........................................... Geological Survey.
   a. White.

   a. Yellowish-white.

   a. Spotted green and white.

In the township of Grenville and its Augmentation, a band of crystalline limestone, which has an extensive run through the country, presents in many places a peculiar variety of marble, having a white ground marked with a number of small green spots, arising from the presence of serpentine; which occasionally forms larger angular masses. These small spots, more or less aggregated, usually run in bands parallel with the bed, and clearly mark the stratified character of the rock. These bands, as in the case of the Arnprior marble, are sometimes even, and sometimes corrugated, giving diversities of pattern in cut surfaces. Sometimes the serpentine, instead of green, is sulphur yellow, as in the specimen from Grenville. In many parts of the country the Laurentian limestones are free from foreign minerals, and give white marbles. These, however, are usually too coarse grained for statuary purposes, and sometimes they are barred with slight differences of color. The specimen from Elzevir, obtained from Mr. Billa Flint, of Belleville, is an instance of this. Many years ago a mill for cutting and polishing a marble like the specimen from the Augmentation of Grenville, was erected on the Calumet, lot 19, range 3, of Grenville, where a similar rock occurs; but the demand for the marble was not sufficient to make the enterprise profitable.—Laurentian.

5. St. Armand ........................................... C. R. Cheeseeman, Phillipsburg.
   a. White.
   b. White.
   c. White, clouded with pale green.
   d. Dove grey, marked with white.

The marbles, of which Mr. Cheeseeman exhibits specimens, occur in great abundance in the immediate vicinity of Phillipsburg, on Lake Champlain. They are all easily cut and take a good polish. Should a railway, which is projected between St. Johns and St. Albans, be carried into operation, it is probable there would be some demand for the stone. No quarries have been opened on any of the beds, and the specimens are taken from surfaces that have long been exposed to the influence of the weather.—Quebec group, Lower Silurian.

   a. Black.

About 1 1/2 miles south eastward from Phillipsburg, there occurs a black marble, similar to this specimen. The beds dip to the eastward at an angle of about 12°, and a quarry was many years ago opened on one of them, which has a considerable thickness. The stone was exported to the United States, and much esteemed in New York, but the opening of quarries of black marble at Glen's Falls, where there is a great water power, interfered with the demand, and caused the enterprise to be abandoned.—Quebec group, Lower Silurian.
38 DESCRIPTIVE CATALOGUE.

   a. Red veined with white.
   This red marble occurs near the river Guillaume, associated with red shales and sandstones, resembling those of Sillery, near Quebec. The red limestone is succeeded by a band of peculiar argillaceous rock resembling the gabro rossi of the Italians.—Quebec group, Lower Silurian.

   a. Grey
   b. Grey with red spots.
   Similar grey marbles, with red spots, occur in the same formation as the rock of Caughnawaga, behind the city of Montreal, and on Île Bizard; while beds in the same formation at St. Lin, in the county of L'Assomption, are wholly red. In all of these localities the rock is filled with fossils, which are plainly seen on the polished surfaces.—Chazy formation, Lower Silurian.

   a. Dove grey.
   The marble of St. Dominique is easily cut, and takes a good polish. It is surprising that situated so near to Montreal, with a railway running near, it has not been applied to various purposes in the city, for which a stone not so good is at present used.—Chazy formation, Lower Silurian.

   a. Grey, with thickly disseminated white spots.
   b. Dark-grey with more thinly disseminated white spots.
   The bed from which the specimen (a) is taken, varies in thickness from 3 to 6 inches; it is near the surface, and easily quarried, but it has hitherto been but little used. The locality is 1 mile from the S. bank of the Ottawa, 4 miles west of L'Orignal village and 64 above Montreal. The white spots are caused by small bivalve shells (Atrypa plena,) filled with calcite. Of the darker variety (b) there are two beds, of 6 inches and 1 foot respectively, near the surface, and overlying the previous bed (a). Blocks large enough for chimney and mantle pieces are readily obtained.

   a. Drab.
   This drab colored marble occurs in great quantity on Esquimaux Island, of the Mingan group, where the stone might be easily loaded on board of small vessels. It cuts with great facility, and takes a uniform polish.—Chazy formation, Lower Silurian.

   a. Brownish black.
   b. Greenish black.
13. **Cornwall** ........................................... *Geological Survey.*

   a. Black.

   These black marbles, from Pointe Claire and Cornwall, are derived from one or two beds about 2 feet thick each, at the base of the Birdseye and Black River formation, and they are apparently the only beds of the formation that take a sufficiently even polish to be fit for the purpose. In the higher beds there are patches, which, from being more argillaceous than other parts, receive but an inferior polish, and produce a bad effect.—*Birdseye and Black River formation, Lower Silurian.*

14. **Pakenham** ........................................... *Geological Survey.*


   The Birdseye and Black River formation at Pakenham, on the Mississippi, a tributary of the Ottawa, yields a very peculiar dark smoke-brown or snuff-brown marble. The stone takes a good polish; but small pieces of chert are sometimes met with, which renders it necessary to be careful in selecting slabs to be wrought. Mr. Dickson, of Pakenham, on whose property the bed occurs, and from whom the specimen exhibited was obtained, had at one time fitted up an apparatus, driven by the waste power of his saw-mill, to polish slabs for chimney pieces and other ornamental purposes. But there was, at that time, no consumption for the material in the neighborhood, and no railway for carriage to a distance, and the marble works were abandoned.—*Birdseye and Black River formation, Lower Silurian.*

15. **Gloucester** ........................................... *Geological Survey.*

   a. Brownish grey.


   a. Grey from the *Trenton* formation.

   b. Grey from the *Chazy* formation.

   The Montreal marble is derived from a bed in the Trenton and another in the Chazy formation. Slabs for chimney pieces and table tops are sawed and polished by Mr. Hammond, and used for common purposes.—*Trenton and Chazy formations, Lower Silurian.*

17. **Dudswell**, lot 23, range 7 ........................................... *Geological Survey.*

   a. Cream white, striped with yellow.

   b. Dark grey and yellowish.

   c. Fawn yellow and white.

   Were the limestones of Dudswell worked, it is probable good marble might be obtained from them. The specimens exhibited, of cream white and yellow, and dark grey and yellow, are from beds that overlie one another. The yellow streaks in both of these marbles are composed of dolomite, while the light ground of the one and the dark ground of the other are carbonate of lime. When the dark grey approaches black, which it sometimes does, and the yellow streaks are narrow, the marble bears a strong resemblance to the Portor marble from Northern Italy, sometimes known as black and gold. On analysis the resemblance between the two is further sustained by the fact, that in both cases the ground is a pure limestone, and the yellow veins are dolomite. It is not unlikely, that if the rock were extensively quarried, some beds might be found in which the resemblance to the Portor would be closer than in the specimens exhibited.—*Upper Herberg formation, Devonian.*
Serpentines.

1. Orford, lot 6, range 13 .............................. Geological Survey.
   a. Brecciated, dark green.
   b. " light green.

2. Orford, lot 12, range 8 ............................... Geological Survey.
   a. Dark green, striped with light green.

   a. Green and white.
   b. Dark and light green.

   a. Brecciated green.

   a. Brecciated green, veined with white.

The band of serpentine, from different places on which, these specimens have been obtained, has been traced on the south side of the St. Lawrence, from Potton to Cranbourne, 140 miles; in 40 miles of which it is repeated twice by undulations, giving an additional 80 miles to its outcrop. It is again recognized 250 miles farther to the N. E., in Mount Albert in the Shickshock Mountains, and about 70 miles farther in Mount Serpentine, approaching Gaspé Bay. All the specimens of these rocks, which have been analysed, contain small quantities of chromium and nickel, and the band is associated in its distribution with soapstone, potstone, dolomite and magnesite. The whole of these occur in large quantities, and in them, as well as in the serpentine, chromic iron occurs, sometimes in workable quantities. These rocks, or others immediately near them, contain the metals iron, lead, zinc, copper, nickel, silver and gold, and with the drift gold, derived from these strata, are platinum, iridosmine, and traces of mercury. In 1847, these serpentines, from their distribution, were described in the reports of the Geological Survey as an altered sedimentary rock. All subsequent observations have confirmed this, and beautifully stratified masses of it have at length been discovered in Mount Albert.—Quebec group, Lower Silurian.

None of the serpentines, and with the few trifling exceptions that have been mentioned, none of the marbles of Canada, have yet been quarried for economic purposes. All of the specimens of them exhibited by the Geological Survey, are consequently from parts of the strata that have long been exposed to the influence of weather, and are of course inferior to the unweathered portions beneath. There appears little doubt that in time both the limestones and serpentines will afford a great amount of beautiful material for architectural adornment, and support a great amount of industry.
SLATES, FLAGSTONES, LIME, BRICKS, AND DRAIN TILES.

Roofing Slates.

1. Walton Quarry, Melbourne, lot 22, range 6. 

- Specimens of slate.

This band of slate is in immediate contact with the summit of the serpentine. It has a breadth of a mile, and dips about S.E. 80°. Mr. Walton commenced opening a quarry upon it in 1860, and found it necessary, in order to gain access to the slate, to make a tunnel through a part of the serpentine. To complete this, and to expose a sufficient face in the slate to pursue profitable working, has required 2 years of time, and $30,000 of expenditure. The face now exposed has a height of 75 feet; but the band of slate crosses the St. Francis, and the fall from the position where the quarry is now worked, to the level of the stream, is upwards of 400 feet, the distance being 1.2 miles; so that by commencing an open cutting on the slate, at the level of the stream, a much greater exposure can be ultimately attained.

Up to a comparatively recent period, the usual coverings of houses in Canada have been wooden shingles, galvanized iron or tinplate, but so many destructive fires have occurred from the use of the first of these, that they are now interdicted in all large towns. Slate, as a covering, costs about $1 more than shingles, but $0.5 less than tin, and $0.3 less than galvanized iron. In the following table are shown, 1st, the sizes of the slates, in inches; 2nd, the number of such slates in a square (of 100 square feet); and, 3rd, the price per square at which Mr. Walton supplies his slates, placed on the railroad cars at Richmond, which is within 1.2 miles of the quarry.

<table>
<thead>
<tr>
<th>Sizes</th>
<th>Number</th>
<th>Price</th>
<th>Sizes</th>
<th>Number</th>
<th>Price</th>
<th>Sizes</th>
<th>Number</th>
<th>Price</th>
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<td>447</td>
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<td>277</td>
<td>3.60</td>
<td>12 x 6</td>
<td>533</td>
<td>2.25</td>
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</table>

The quarry has now been in operation since the spring of 1861; 2000 squares have been sold, and some of the slates have been sent to a distance of 250 miles from the quarry; a quantity of them having been purchased for Sarnia on the River St. Clair. To show that slate as a covering is well adapted to resist the influences of a Canadian climate, it may be here stated that slates from Angers in France, have been exposed on the roof of the Seminary building on the corner of Notre Dame and St. François Xavier Streets, in Montreal, for upwards of 100 years, without any perceptible deterioration. The strong resemblance between those and the slates of Melbourne, as well as those from Bangor in Wales, may be seen in the following comparative analyses by Mr. T. Sterry Hunt:

<table>
<thead>
<tr>
<th>Welsh</th>
<th>French</th>
<th>Melbourne</th>
</tr>
</thead>
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<tr>
<td>Silica</td>
<td>60.50</td>
<td>67.00</td>
</tr>
<tr>
<td>Alumina</td>
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<td>20.10</td>
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<tr>
<td>Protoxyd of Iron</td>
<td>7.83</td>
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<tr>
<td>Lime</td>
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<td>3.39</td>
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<tr>
<td>Potash</td>
<td>3.18</td>
<td>1.73</td>
</tr>
<tr>
<td>Soda</td>
<td>2.20</td>
<td>1.50</td>
</tr>
<tr>
<td>Water</td>
<td>3.30</td>
<td>4.60</td>
</tr>
</tbody>
</table>

100.03 | 100.13 | 99.63 |

The proximity of the serpentine leaves no doubt as to the geological horizon of these slates.—Quebec group, Lower Silurian.
   a. Specimens of slate.

   a. Specimens of slate.

   a. Specimens of slate.

   a. Specimens of slate.

The Cleveland slates are a continuation of the Melbourne band. The Shipton Slate Company opened a quarry on them in 1854, and found them to be of superior quality. This quarry is now for sale. The slates of Orford may be on the same band, about 10 or 12 miles to S. E.; but the geological horizon of the Tring slates is uncertain, though they probably belong to the Quebec group. The Kingsley slates appear to be lower in the series than the magnesian group of strata.—Quebec group, Lower Silurian.

Flagstones.

   a. Specimens of the flagstone.

This is a hard, fine-grained sandstone; and the surfaces are even and parallel. Many of the beds of the band, which is 20 feet thick, can be split into flagstones; which are used in the city of Toronto. Similar flagstones, used at Hamilton, are obtained from the same band there, and an equally good quality can be obtained wherever the band occurs.—Grey band, Medina formation, Lower Silurian.

Hydraulic lime.

   a. Raw stone,
   b. Prepared cement.

The bed which yields the Thorold cement is a dark brown dolomite of the Clinton formation. During the construction of various railway and other public works the quantity of cement manufactured by Mr. Brown averaged 80,000 bushels annually, but at present the quantity made does not exceed one-tenth of the amount. The present price of the cement is from 20 to 26 cents per bushel of 60 lbs.—Clinton formation, Middle Silurian.

   a. Raw stone.
   b. Prepared cement.

The beds of this deposit are from 2 to 11 inches thick, occasionally separated by layers of shale, making in all 15 feet. Cement has not yet been manufactured from this stone; and none is made within 100 miles of the locality, although there would, no doubt, be considerable demand for it in the neighborhood, were it prepared at the place. The locality is in the bank at a mill-dam on the Saugeen River, where an unlimited water-power for grinding the cement could be had.—Onondaga formation, Upper Silurian.
   a. Raw stone.
   b. Prepared cement.

   This stone occurs in a band of 9 feet thick, in beds varying from 3 to 7 inches. The cement is manufactured in considerable quantities by Messrs. Bescoby and Newton. It sets slowly, and hardens during several weeks, after which it is said to possess great strength.—Clinton group, Middle Silurian.

   a. Raw stone.

   Though the rock occurs in Nepean, its product is usually designated as the Hull cement, from having been manufactured for several years, by Mr. Wright of Hull, opposite to Ottawa. The rock is a limestone holding about 12 per cent. of carbonate of magnesia, and it yields a strong and lasting cement. The bed to which it belongs, has been traced for nearly 100 miles through the country, preserving a very uniform character.—Chazy formation, Lower Silurian.

   a. Raw stone.

   This specimen comes from a band 3½ feet thick, associated with a band of chert, and separating into beds averaging 6 inches. It is not worked, but could be easily quarried, and a good water-power for grinding is ready at the spot.—Niagara group, Middle Silurian.

   a. Raw cement.

   These specimens of black dolomite are derived from the Mountain Portage, about 5 miles up the Magdalen River from its mouth. The stone occurs in beds of from 2 to 4 inches, interstratified in black graptolitic shales, and yields a very strong hydraulic cement, setting in a few minutes under water, to a very hard and tenacious mass of a yellowish color. Similar bands occur at the Grande Coupe, 6 miles below Great River. The range of the formation containing these bands, being from Gaspé to Quebec, makes it probable that a considerable quantity of the stone may be obtained from various places along the south shore of the St. Lawrence. The stone differs from that at Quebec, from which Capt., now Major-General Babcock, R.E., first prepared a cement, now manufactured by Mr. P. Guvrean. This contains no magnesia, while the Gaspé stone is a dolomite.—Hudson River formation, Lower Silurian.

Common lime.

   a. Raw stone.
   b. Prepared lime.

   This lime is burnt from the Guelph magnesian limestone; the stone takes rather longer to calcine than pure limestone; it slacks without the evolution of much heat, to a very white powder, much prized for whitewash and for mortar, which sets quickly. The stone occurs in unlimited quantities.—Guelph formation, Middle Silurian.
   a. Raw stone.
   b. Prepared lime.
   This remarkably white lime is burnt from a band of drab-colored magnesian limestone, 7 feet thick. It makes a superior whitewash and a strong cement. — Onondaga formation, Upper Silurian.

   a. Raw stone.
   b. Prepared lime.
   The limestone which yields the best stone for the purposes of construction at Montreal, also burns to excellent lime, and the débris which accumulates in the process of quarrying the building stone, is used for that purpose. The quantity of lime manufactured at Montreal is estimated to be 270,000 bushels per annum, and the price about $0.16 per bushel.

Common bricks.

   a. Red bricks.
   These bricks are made from a drab-colored clay, which has been dug to a depth of 4 feet. White bricks are made from the same clay by using a different sand. The deposit is not extensive. — Drift.

   a. Red bricks.
   These bricks are made from a bed of 9 feet of purplish-brown finely laminated clay, reposing on 20 feet of highly calcareous sand. — Drift.

   a. Red bricks.
   These specimens are manufactured from a thinly laminated blue clay, which the brick-makers of the place state to be upwards of 100 feet thick, and which requires a mixture of sand for the manufacture. In 1852 about 2,000,000 bricks were manufactured by 7 brick makers. — Drift.

4. Montreal ........................................ Peel & Compte, Montreal,
   a. Common building bricks; price $5 per 1000.
   Messrs. Peel & Compte manufacture 6,000,000 common bricks annually, which are sold at from $5 to $6 per 1000.
   The red bricks of Montreal are manufactured from a blue clay of marine origin, which is interstratified with reddish layers, and runs under a deposit of sand. The clay has been excavated to a depth of 20 feet, and may be deeper, as the same formation is known to
have a greater thickness in other localities. Its marine origin is proved by the occurrence of sea shells, of about 6 species in the pure clay, and about 30 in the sandy clay immediately overlying it; all probably the same as species now inhabiting the ocean. Our knowledge of the fossils of these deposits has been greatly extended by the researches of Dr. Dawson, of McGill College, who has more than doubled the number of shells known a few years since, and added to the list many species of Bryozoa, Foraminifera, and other small forms. The remains of the capelings (Mallotus villosus) and the lump-sucker (Cyclostomus lumpus) are obtained from the same clays near Ottawa, and a clay-pit of Messrs. Peel & Compte, on Côteau Baron, has yielded 19 of the caudal vertebral joints of a cetacean, similar to a species discovered in Vermont by the late Mr. Zadock Thompson, and named by Mr. C. H. Hítecheock, Beluga vermontana. On Côteau Baron these remains were accompanied by one of the pelvic bones of a seal, by sea shells, and by fragments of white cedar, Thuya occidentalis. The locality is about 140 feet above the level of the sea. In another of Messrs. Peel & Compte's pits there has recently been found a nearly entire skeleton of the Greenland seal (Phoca groenlandica,) a species still living in the Gulf of St. Lawrence; from the size of the head, the animal appears to have been 6 feet long and full grown. Within a few days, a clay-pit of Messrs. Bulmer and Sheppard has given many of the bones of some other animal, supposed to be a seal, of much smaller dimensions. The brick yards are situated to the north-west of Mount Royal, on a plateau of considerable extent, above which, well-marked sea margins occur on the sides of the mountain, at elevations of 223, 386, 440 and 570 feet above the sea level, with marine shells up to the last mentioned height.—Allison.

   a. Common building bricks, price $5 per 1000.
   b. Pressed front bricks, "$ 1.00"
   c. Radiating front bricks "$ 1.75"
   d. Circular bricks for shafts "$ 1.25"

The quantity of bricks manufactured by Messrs. Bulmer & Sheppard is equal to 6,000,000 per annum. In this manufacture they use Boaden's brick-making machine.—Allison.

   a. White bricks.

   The specimens are manufactured from a brownish laminated clay, which burns white, and underlain by a considerable deposit of sand. Either red or white bricks are made of this clay, according to the sand used.—Drift.

   a. White bricks.

   The deposit of clay, from which these white bricks are manufactured at Toronto, has a thickness exceeding 20 feet, and extends eastward at least as far as Cobourg. It appears to be unconformably overlaid by a bed, which is 3 feet thick, giving red bricks, for while the white brick-clay lies in very even horizontal strata, the other undulates with the general surface, not however descending to the bottom of deep ravines. The average annual manufacture of white bricks in Toronto is from 3 to 5 millions, and the ordinary price at the kiln is from $5.50 to $6 per 1000. The price of common red bricks is from $3 to $4 per 1000, and the average annual manufacture, including all kinds, is from 8 to 10 millions.—Drift.
Drain tiles.

1. North Plantagenet........................................... C. P. Treadwell, L’Original.
   a. 1½ inch red agricultural drain tiles.
   These tiles are manufactured by Thomas Gibb, at Treadwell, North Plantagenet, from a blue clay which forms a considerable deposit on the banks of the Ottawa. The price of these tiles is $10 per 1000.

2. Quebec .................................................... H. O’Donnell, C.E., Quebec.
   a. Clay used in making sewerage pipe tiles.
   b. A 6-inch sewerage pipe tile.
   These tiles are manufactured by Messrs. W. & D. Bell, from a deposit of clay, varying in thickness from 3 feet to 30 feet, on the river St. Charles, between 1 and 2 miles from Quebec. They are used for main sewers and house drains, in the city of Quebec, where 161,000 feet of them have been laid. They are united by means of rings of the same material, which cover the joints, and permit alterations and repairs without breaking the pipes. When in place, the pipes are capable of resisting a pressure of 50 lbs. to the square inch, and, when properly glazed in the inside, they remain free from the incrustations which are found to gather on the inside of iron pipes. The prices of these drain-tiles are:

<table>
<thead>
<tr>
<th>Diameter (inches)</th>
<th>Price per linear foot</th>
</tr>
</thead>
<tbody>
<tr>
<td>4 in.</td>
<td>$0.15</td>
</tr>
<tr>
<td>6 in.</td>
<td>$0.21</td>
</tr>
<tr>
<td>9 in.</td>
<td>$0.33</td>
</tr>
<tr>
<td>12 in.</td>
<td>$0.60</td>
</tr>
<tr>
<td>15 in.</td>
<td>$0.86</td>
</tr>
<tr>
<td>18 in.</td>
<td>$1.13</td>
</tr>
</tbody>
</table>

GRINDING AND POLISHING MINERALS.

Whetstones.

1. Stanstead, lot 15, range 1................................ Geologic Survey.
   a. Cut whetstones.

   a. Cut whetstones.

   a. Cut whetstones.

4. Kingsey, lot 7, range 2.......................................... Geologic Survey.
   a. Cut whetstones.

In the Eastern Townships, stones of a good grit for the purpose of whetstones are found in several places. A band of rock runs from Whetstone Island in Memphremagog Lake, lot 15, range 1, of Stanstead, by Lee’s Pond to the head of Massawippi Lake, in Hatley; a distance of nearly 12 miles, and the band may be available much further. The rock appears to be a mica slate, passing into an argillite, and its stratigraphical place would seem to be above the magnesian series. There is also a range of whetstone rock on each side of the anticlinal running from Melbourne to Danville, beneath the magnesian rocks. This rock again appears on the north-west side of the Shipton and St. Armand synclinal, in Kingsey, and good samples of the stone occur on lot 7, range 2 of the township, where whetstones were some years ago manufactured by Messrs. Gilmour & Jackman. They are much softer than the Memphramagog stones, the rock being probably more argillaceous. The Bolton stone very much resembles that of Memphramagog, but its stratigraphical place is probably the same as that of Kingsey.—Quebec group, Lower Silurian.
   a. Cut specimens.

   These whetstones are obtained from about 20 feet of thin, even bedded, and very fine
grain sandstones and arenaceous shales, at the top of the Hudson River formation. The
inhabitants of the neighborhood make whetstones for their own use, from this rock, but
it has never been extensively worked. The same rock is found in the same geological posi-
tion at Meaford, Cape Rich, and on the Grand Manitoulin Island.—Hudson River forma-
tion, Lower Silurian.

   a. Cut specimens.

   The specimens are taken from about 20 feet of freestone, representing the Grey band.
The rock is in every way suited to make superior scythe stones, although they have never
yet been manufactured from it.—Medina formation, Middle Silurian.

   a. Cut specimens.

   These specimens are from a few feet of very fine grained compact sandstone at the foot
of the falls, and immediately underlying the dolomite of the Clinton formation. It ap-
ppears to be the upper part of the Grey band. The rock is not worked in this locality.—
Medina formation, Middle Silurian.

   a. Cut specimens.

   The mica slates associated with the crystalline limestones of the Laurentian series are fre-
cently of the character required for scythe stones, and a band of this description occurs
in Madoc, on the property of Mr. O'Hara, who at one time cut and wrought the rock into
whetstones for sale. The whetstone rock occurs not far from crystalline limestone, and in
immediate contact with a thick band of conglomerate, of which the matrix weathers white,
and appears to be a dolomite. The pebbles, which are frequently large, some of them
being 6 inches in diameter, are chiefly of quartz, but there are others of feldspar, and some
which are calcareous. The quartz pebbles are for the most part distinctly rounded,
and their colors various, some being bluish, and others white or pinkish on fracture.
Those of feldspar are red and white.—Laurentian.

Hones.

   a. Cut specimens.

   Some of the silicious slates of the Huronian series yield very fine ones. They are
usually of a green color, and occupy a place in the lower part of the series.—Huronian.
Grindstones.

1. Nottawasaga, lot 24, range 11

This grindstone, which measures 2 feet 4 inches in diameter, is from the Grey band, which is about 20 feet thick at this locality, and the whole of it appears equally well qualified for making grindstones. It splits well into the various thicknesses required for these stones, and they have been made from it, by hand, in considerable numbers, both at this place, and in the township of Mulmur. The same rock is found in many places near the escarpment of the Niagara formation, in Nottawasaga and Mulmur. The grindstones made from it are declared by practical men to be superior to those imported; but they have never yet been manufactured by machinery. A lathe for turning them could be erected on one of the numerous streams which cross the formation, for about $1000 (£200 stg.). Grindstones roughly hewn by hand, sell for 1½ cents per pound on the spot, which is the price of the imported Ohio stones, as sold on the coast of Lake Huron.—Grey band, Medina formation, Middle Silurian.

Millstones.

1. Grenville, lot 3, range 5

This buhrstone occurs on the property of Mr. James Lowe. On his land, and that of some of his neighbors, it constitutes a series of veins, cutting an intrusive mass of syenite, which occupies an area of 36 square miles, among the Laurentian rocks of Grenville, Chatham, and Wentworth. The veins consist of yellowish-brown or flesh-red cellular chert; the colors, in some cases, running in bands parallel to one another, and sometimes being rather confusedly mingled, giving the aspect of a breccia. The cells are unequally distributed, some parts being nearly destitute of them, while in others they are very abundant and of various sizes, from that of a pin's head to an inch in diameter. On the walls of some of the cells, small transparent crystals of quartz are implanted; and in some of them are impressions of cubical forms, resulting, probably, from crystals of fluor spar, which have disappeared. The stone has the chemical composition of flint or chalcedony. On Mr. Lowe's ground, one of the veins runs nearly east and west, and stands in a vertical attitude; while its breadth varies from about 4 to about 7 feet. When the vein is banded, the colors run parallel with the sides. The attitude and associations of the chert clearly show that it cannot be of sedimentary origin, and its composition, taken in conjunction with the igneous character of the district, suggests the probability that it is an aqueous deposit, which has filled up fissures in the syenite, and is similar in its origin to the agates and chalcedony, which in smaller masses are common in various rocks. For a distance of perhaps 300 yards on each side of these veins of chert, while the quartz of the syenite remains unchanged, the feldspar has been more or less decomposed, and is converted into a sort of kaolin. As this process involves a separation of silica from the feldspar, it is not improbable that it has been the origin of the veins of silex.—Laurentian.
MINERAL MANURES.

Gypsum.

1. Oneida ...................................................... Thomas Martindale.
   a. Crude.
   b. Prepared.

2. Oneida ..................................................... I. Donaldson.
   a. Crude.
   b. Prepared.

   a. Crude.
   b. Prepared.
   c. Plan of the mine by Mr. J. De Cew.

All the gypsum mines at present worked in Canada, occur on the Grand River, in a distance of 36 miles, extending from Cayuga to Paris. The formation, to which they belong, however, runs from the Niagara River to the Saugeen, on Lake Huron, a distance of about 150 miles; and it seems probable that as the country to the north-west of Paris becomes more settled, further discoveries of workable masses will be made in that direction. All the mines appear to be confined to one stratigraphical position in the formation, which is probably about the middle. The mineral occurs in lenticular masses, varying in horizontal diameter, from a few yards to a quarter of a mile, with a thickness of from 3 to 7 feet. The layer of gypsum appears to be in general both underlaid and overlaid by beds of dolomite, much of which is fit for the purposes of hydraulic cement, and the gypsum itself is sometimes interstratified with thin beds of dolomite. In some parts, there appear to be two workable ranges of gypsum, one a few feet above the other. But this probably, is only to be considered a thickening of the gypsiferous band, with an interstratification of a larger mass of dolomite.—Onondaga formation, Upper Silurian.

Mr. Taylor, at his present rate of working, obtains annually from his mine about 4000 tons of gypsum, and the following are the prices:

- Plaster in rock, .............................................. $2.00 per ton.
- " ground for land purposes, ................. 3.50 "
- " " for stucco, raw, ...................... 7.00 "
- " " calcined ......................... 10.00 "

Fresh-water Shell Marl.

   a. Specimen of the marl.

This deposit is on the property of Messrs. John & Thomas MacKay, of Rideau Hall, New Edinburgh, and is 5 feet thick. Among the shells which it contains, the following species are met with: Plana heterostrophia, Limono pallida, Planorbis bicarinatus, P. campyloides, P. parvus, Amnicola porata, and Tricaria tricarinata. With a thin covering of vegetable mould, the marl supports a growth of large forest trees, under which it extends some distance along the east side of a small lake or pond, which occurs in the course of a small stream, discharging by a narrow ravine into the Ottawa close by. The surface of the pond...
is 26 feet above the river in summer, but only 8 feet in the freshets of spring; the river in summer is 118 feet above the sea. The marl bed is on a level surface, 26 feet above the pond, and, after spreading over a breadth of 200 yards, it appears to run under a terrace five feet higher, which maintains a level surface for considerable distance. This, instead of overlying the marl, may be the margin of the lake in which it was deposited. The pond is 300 yards wide, and on the west side there are evidences of three periods of recession, in distinct terraces; which are at heights of 30, 60 and 75 feet, respectively, over the level of the pond, or 174, 204 and 219 feet above the sea, each with a sudden step rising to the next. The upper step, or perhaps the upper two steps, may exhibit former limits of the sea. The clays of the banks of the Ottawa, at this part, are of marine origin, and 9 miles farther down the river, at Green's Creek, hold the remains of two species of sea fish, which have been already mentioned (page 45), the Mallotus villosus, or common capeling, and the Cyclostoma lumpus, or lump-sucker; with Scrobiera rugosa, Leda Portlandica, and other sea shells. The two flappers of a seal were obtained from the same clay, as well as sea-weeds, and leaves of large exogenous trees.—Alluvion.

2. Sheffield, lots 15 and 16, range 2

a. Specimen of the marl.

This deposit, which is on the property of Mr. McConnel, extends over an area of 200 acres and perhaps more, with a thickness, over the greater portion, of at least 10 feet. On the surface there is a thin soil, bearing a luxuriant growth of prairie grass. The species of shells observed here are Planorbis bicarinatus, P. parvus, Physa heterostropha, Amnicola porata, with undetermined species of Limnana, Valvata, Cyclas, and Pisidium. Another locality in Sheffield, where marl occurs, is on lot 12, ranges 3 and 4, extending over at least 300 acres and perhaps more than 400. The place where it occurs is chiefly a swamp or marsh, and it is covered over by an accumulation of excellent peat, averaging 4 feet in thickness. Still another locality in the same township, is in White Lake, and the brook leading from this to Beaver Lake.—Alluvion.

3. Montreal

a. Specimen of the marl.

This deposit, which is very pure and white, occurs at Thornberry on the west side of Mount Royal. It is overlaid by peat, but it does not seem to be of very great extent. The species of shells met with in it are Planorbis campanulatus, P. bicarinatus, P. triovulvis, P. parvus, Limnana umbra, L. stagnalis, Physa marginata, P. heterostropha, Valvata bicornata, Amnicola porata, Melania acuta, Cyclas similis, Pisidium dubium, and an undetermined Unio.—Alluvion.

4. Nepean

a. Specimen of the marl.

This deposit is on the property of Mr. Sparks, of Ottawa. It is a foot thick, and is covered with a thin layer of peat. The species of shells found in it are Physa heterostropha, P. marginata, Planorbis bicornatus, P. parvus, P. campanulatus, Limnana modestella, Amnicola porata, Valvata tricarinata, and Pisidium.—Alluvion.
   a. Specimen of the marl.

   The marl is found on the property of Mr. George Cross, in the bottom of a prairie-like flat, traversed by a small brook; it is known to cover between 3 and 4 acres on this lot, but it is believed to be more extensive, and to continue into the next lot eastward. The specimen was obtained near the edge of the deposit. The bed is here 32 feet deep, and is overlaid by 4 feet of peat. The surface is overgrown with grass, reeds, and moss, and the locality appears to have been the former site of a small lake. The marl taken from the upper half of the bed becomes white when dry, and is filled with well preserved shells; that from the lower half is of a bluish color and more tenacious character. Branches and trunks of trees, in a good state of preservation, are found in the marl, but not in the peat. The marl has proved a very efficacious manure to the adjoining lands, which are of a sandy character. In digging it the effluvium evolved is so offensive, that few men can bear it. The peat is also used as a manure by the proprietor. The following species of fresh-water shells have been obtained from this marl: Limnaea stagnalis, L. umbrosa, Planorbis tricostatus, P. campanulatus, P. bicarinatus, P. percaria, Physa heterostropha, Amnicola porata, Valvata tricarinata, Cyclas similis, and an Anodonta.—Alluvion.

   a. Specimen of the marl.

   The marl here occurs in a flat meadow, skirting a small stream, and extends over an area of 7 acres. The bed is 2 feet deep, and is covered by a foot of peat, which supports a growth of prairie grass. The marl from the lower part of the bed is of a blue color when wet, while that from the middle is whitish and has been used by the people of the neighborhood as a whitewash, but not yet as a manure, the lands being naturally very calcareous. Most of the shells are finely comminuted, and only an occasional whole specimen preserved. These appear to belong altogether to small species, and among them occur Planorbis percaria, Valvata humeralis, V. tricarinata, Amnicola porata, and several small species of Pisidium.—Alluvion.

   a. Specimen of the marl.

   This deposit is about 6 acres in extent, with an uncertain depth of 21 feet. It is very white, and overlaid by a thin stratum of black mould. The surface has the aspect of prairie land, and is intersected by a brook. Similar prairies, in which marl is said to be found, occur at intervals along the brook, for 4 miles, and the whole area underlaid by marl, is estimated at 40 acres. It has hitherto been used only for whitewashing. Among the shells which it contains, the genera Limnaea, Planorbis, Physa, Valvata, Amnicola, Cyclas and Pisidium are represented. A great many deposits of marl, similar to this and to the last, are met with in the counties of Bruce and Grey.—Alluvion.

   a. Specimen of the marl.

   This bed occurs in low ground, close to the town of Durham. Its extent is uncertain, but it is known to cover 8 or 10 acres. At the spot where the specimen was taken, its depth was 4 feet. It is very solid and pure, and is covered by heavy timber. Physa heterostropha, Planorbis percaria, Valvata tricarinata, Amnicola porata, with small species of Pisidium, are among the shells which it contains.—Alluvion.

a. Specimen of the marl.

Marl Lake, at the west end of Anticosti, has a superficies of about 90 acres, and appears to have a bottom covered with shell marl. The thickness of the marl seems to be considerable, but its exact measure has not been ascertained. The creek which empties the lake into Indian Cove, carries down a large quantity of the marl to the sea, where it becomes spread out for a considerable distance over the rocks of the vicinity. This is the most northern deposit of marl which has been met with. Among the species of shells which it contains are Limnea acuta, Planorbis parvus, P. trivolvis, with another small undetermined species, Physa heterostropha, Valvata sincera, Psidium dubium, and one or two species of the last genus, supposed to be new. The most abundant species observed is Limnea acuta, (Lea,) the next most abundant is Valvata sincera, (Say,) Two small species of land snails were met with in the marl, Helix arborea and H. striatella.—Alluvion.


a. Specimen of the marl.

This deposit is on the land of Mr. Yeoman of Belleville, but does not appear to be extensive. The species of shells observed are Valvata humeralis, Psidium dubium, with an undetermined Limnea and a Psidium.—Alluvion.


a. Specimen of the marl.

This shell marl occurs on a pond, a mile south-east of Phillipsburgh, on lots 156 and 157 of St. Armand, on the lands of Mr. Strite and Mr. Taylor. The marl is visible all around the pond, and consists of the comminuted remains of fresh water shells to a depth of several feet, resting on a deposit holding marine shells, probably of the age of the drift. The fresh water species are Planorbis parvus, P. campanulatus, Limnea umbrosa, Physa heterostropha, Valvata tricarinata, and Amnicola porata. The whole depth is in some parts 7 feet, and the area of the deposit may be between 30 and 40 acres. The specimen exhibited was obtained from Mr. Strite.—Alluvion.

Calcereous Tufa.

1. Noisy River Falls ..................................... Geological Survey.

a. Specimen of the tufa.

This tufa covers the extensive slopes on both sides of the river, from the base of the Niagara escarpment to the edge of the water. It is constantly soft and moist, and is cut into by numerous springs, which flow down the long slopes. It probably covers an area of 300 acres in the vicinity of the falls, with an average thickness of 5 feet. Tufa of this character is found in many places along the base of the Niagara formation, in the counties of Grey and Simcoe; the most important is that on the great slopes of the Beaver River, in Euphrasia and Artemesia, which is supposed to extend over more than 1000 acres, in the form of a strip on each side of the river.—Alluvion.
7.

MINERAL PAINTS.

Iron ochres.

   a. Brownish ochre.
   b. Brownish-black ochre.
   c. Yellow ochre.

   This deposit of ochre is situated on the property of Mr. E. Caron, about a mile and a
   quarter above the mouth of the Ste. Anne River. It appears to extend over about 4 square
   acres. The locality is on the top of a bank, overlooking the main road, from which it is
   removed about a quarter of a mile. The surface of the bed has a slope to the south-east,
   of about 50 feet in about 150 yards, but its bottom keeps nearly level with the lower side
   for some distance back, and then rises quickly to the higher side. The thickness of the
   deposit is thus 17 feet in the deepest part, and varies from that to 4 feet. Its form gives
   great facilities for excavating the ochre, as by beginning on the lower side, a considerable
   face of it would be exposed, and the water would run from it, without the necessity of cut­
   ting drains. The three colors exhibited occur at the surface, but the lower and by far the
   larger part, is of a pale green color. In this green portion the iron is in a lower state
   of oxidation than in the yellow, but like it becomes red upon ignition in the air.—
   
   Alluvion.

   a. Greenish-black ochre.
   b. Yellow ochre.

   In the St. Malo range of the seigniory of Cap de la Madelaine, about 2 miles below the
   church, and 2 miles back from the St. Lawrence, there is a deposit of ochre, extending
   over about 400 square acres. It is interstratified by peat, and underlaid by shell marl,
   which in successive borings along a transverse section from S. E. to N. W., were found
   to be arranged as follows, in descending order,—ochre, peat, and marl being indicated
   by the letters O, P, M:—

   | Paces, 50 | 100 | 145 | 181 | 231 | 441 |
   | Ft. in.   | Ft. in. | Ft. in. | Ft. in. | Ft. in. | Ft. in. |
   | O, 0 6   | O, 2 0   | O, 1 6   | O, 2 0   | P, 9 0   | O, 2 0   |
   | P, 0 6   | P, 4 0   | P, 8 0   | P, 4 0   | M, 0 6   | P, 7 0   |
   | O, 0 6   | P, 2 0   | O, 9 0   | P, 6 0   | M, 0 6   | P, 7 0   |

   In the remaining 290 paces, the ochre is wanting, and we have 12 feet of peat, gradually
   thinning out. A very great quantity of red and yellow ochres might be obtained from
   this locality, and where the ochre is mixed with the peat, masses of the mixture might
   be cut out and dried, and afterwards burned. Experiments on a small scale shew that
   the quantity of peat in the mixture is often sufficient to calcine the ochre.—Alluvion.
3. **Pointe du Lac**

*Geological Survey.*

- **a. Purplish ochre.**
- **b. Yellow ochre.**

An ochre bed of about 400 acres in extent, is situated on the St. Nicholas range of Pointe du Lac Seigniory, on the property of Mr. Pierre Chaillon and his brother. Its thickness varies from 6 inches to 4 feet, and it may have an average of about 18 inches. Its prevailing colors are red and yellow, but there occurs also in some parts a beautiful purple ochre, and in others a blackish-brown. In 1851, Messrs. H. A. Monroe & Co., of New York, made arrangements to prepare the ochres for sale. Two furnaces were erected in the vicinity of the ochre bed, and an agent established to carry out the details of the manufacture, and to attend to the forwarding of the prepared ochre to New York; where the sale of it was effected. From the natural tints that have been mentioned, eight colors are said to have been prepared. The deposit being but little mixed with sand, the chief impurities to be got rid of consisted of the roots of those plants which had been growing on the surface; some of which were found to penetrate to a considerable depth. Two modes were resorted to for this purpose; one by dry sifting, which was used when the natural colors of the ochres were to be preserved, as in the case of the yellow, the purple, and the blackish-brown varieties. The other mode was by burning. The yellow is a hydrated peroxyd of iron, the purple also is probably in some peculiar state of hydratation, but the red is the anhydrous peroxyd. By exposure to a sufficient heat, the water of combination is driven off from the yellow and purple, and both becoming anhydrous peroxyd, assume the tint of the natural red ochre, from which, as from the other two, the vegetable matter in this operation is burnt out. The blackish-brown variety is scarcer than the others, and affords colors of a more valuable description. Purified from roots, without fire, it is sold under the name of raw sienna; and is admirably adapted for graining. When subjected to fire, it assumes a brown of less intensity, and is sold as burnt sienna. As it does not turn red by burning, it is probable that there may be in this ochre an admixture of manganese. The enterprise at Pointe du Lac appears for the present to be abandoned.—*Alluvion.*

4. **Nottawasaga, lot 2, range 11**

*Geological Survey.*

- **a. Yellow ochre.**

This deposit covers about half an acre, on the south bank of the river, and is produced by chalybeate springs issuing from the Clinton formation. When dry, it has a good yellow color. An excavation of 2 feet, near the centre of the deposit, did not reach the bottom. Small deposits of yellow ochre are met with in similar situations near the Clinton formation in other places.—*Alluvion.*

Owen Sound, town plot

*Geological Survey.*

- **a. Yellow ochre.**

This ochre contains a small amount of calcareous tufa, but is of a bright yellow color. The bed occurs at the foot of a bank, in which the Clinton formation crops out, on the S. W. side of the town; its extent has not been accurately ascertained, but it does not seem to be great. It appears to have been deposited by springs which have long since changed their course, and is 4 feet deep in the middle, thinning out towards the edges.—*Alluvion.*
MINERALS OF CANADA.

Sulphate of Barytes.

   a. Specimen from a vein.

   a. Specimen from a vein.

The barytes of Burgess and Lansdowne is derived from veins intersecting Laurentian rocks. In the latter township, as well as in Bedford, the mineral, associated with calc spar, constitutes the veinstone of some of the lead lodes met with there. The vein yielding the Lansdowne specimen cuts Laurentian limestone. In an unsuccessful attempt to mine the vein for lead, it was ascertained that 28 feet of the lode, with a breadth of 27 inches, consisted of highly crystalline almost colorless barytes, of which the vein in this part would yield about 10 tons to a square fathom in the plane of the lode. The most abundant source of barytes in Canada, so far as known, appears to be the veinstones of lodes carrying copper ore, on the north shore of Lake Superior, between Pigeon River and Fort William, and in Thunder Bay. These, however, belong to the Quebec group. In Canada the mineral is as yet applied to no use, but in some parts of the United States it is refined and ground in large quantities, for use as a paint. The value of the crude barytes suited for such a purpose, is about $10 per ton, while the wholesale price of the paint is $30 per ton.

Laurentian.

MINERALS APPLICABLE TO THE FINE ARTS.

Lithographic stone.

   a. Prepared specimen, with fac simile autographs of Canadian Governors.

At Marmora, the Laurentian rocks are overlaid by about 20 feet of brownish-grey and light brownish-buff un fossiliferous compact limestone, with a conchoidal fracture, several beds of which would be well suited for the purposes of lithography, were it not for small imbedded lenticular crystals of calcareous spar, which, when abundant, unfit the stone for such an application. One of the beds, however, which is 2 feet thick, and of impalpable grain, is a lithographic stone of excellent quality. The lower half is much better than the upper, which is somewhat affected by the lenticular crystals of calc spar. The upper inch, which is just above the thus marked part, fits upon it in tooth-like projections, having columnar sides at right angles to the bed, of an inch long in some places; and usually covered with a thin film of bituminous shale. The same tooth-like forms occur in the lower part, but they are there more obscure. The band to which the bed belongs, presents occasional exposures of a similar character, all the way from Hungerford to Rama, a distance of 100 miles; but though the stone has been highly commended by all the lithographers who have tried it, no one has attempted to quarry it for use. The stone exhibited, presents the fac simile autographs of all the governors of Canada, both French and English, from the time of Champlain in 1612 to that of Lord Monck in 1862; with the exception of two of the French governors in the first half of the seventeenth century.-Birdseye and Black River formation, Lower Silurian.
2. Brant, lot 31, ranges 1, 2 .......................... Geological Survey.
   a. Prepared specimen, with Bank cheque and transfer, No. 1.
   b. " " " " No. 2.
   c. " " showing natural fracture, with vignette of an Indian chief.

These are specimens of magnesian limestone of a yellowish drab color and fine texture, with a conchoidal fracture. The locality is the bed of a small stream, on lot 31, between ranges 1 and 2, south of the Durham road, Brant, and about 1 1/2 miles south of the village of Walkerton. About 16 beds of stone, apparently of the same character as the specimens, occur in a vertical section of 9 feet, the thickest being 11 inches. Layers of dark colored shales separate some of the beds. The band is underlain by about 66 feet of soft clayey strata, constituting the bank of the Saugeen River, at the top of which it occurs. The existence of this stone being a very recent discovery, only a preliminary trial of it has been made. The beds from which the specimens are taken are intersected by a number of parallel joints, which render the specimens procured somewhat narrow; but the geological place of the band having been ascertained, it is probable that wider slabs may be found on the strike, in some other locality.—Onondaga formation, Upper Silurian.

   a. Prepared specimen. with drawing of a steam ship.
   b. Transfer in two colors from a.

This stone is of the same character and from the same formation as the last. The locality is at the edge of the river, on the east side of the lot indicated in Brant. Two beds, of 4 and 6 inches respectively, occur here, but they were covered with water at the time the place was visited.—Onondaga formation, Upper Silurian.

MINERALS APPLICABLE TO JEWELLERY.

Agates.

   a. Specimens cut and polished.

These agates occur on the south and north shores of Lake Superior, particularly on the island of St. Ignace, and on Simpson's Island to the east of it; but the largest and best are derived from the trap of Michipicoten Island, where they strew the shore in great abundance. On this island, agate occurs not only in the form of nodules in the trap, but in veins, filling cracks and dislocations, which traverse the trap, and run in several directions.
—Quebec group, Lower Silurian.

Labradorite.

   a. Cut and polished specimens from boulders.
   a. Cut and polished specimens from a bed.

   This beautiful opalescent mineral occurs in disseminated cleavable masses, imbedded in a
cleaner-grained paste of the same mineral character, but destitute of opalescence. The rocks
composed of the series of triclinic feldspars, to which this mineral belongs, have been termed
anorthosites, in describing the Laurentian system; where they occupy a very conspicuous
place. Great mountain masses of the rock occur in Abercrombie, in the county of
Terrebonne, and boulders derived from them lie scattered over the plains to the south.
They are abundant in the neighborhood of Grenville, on the Ottawa.—Laurentian.

Albite (peristerite).

   a. Specimens cut and polished.

   This mineral, the peristerite of Thompson, so called from its beautiful bluish opales-
cence, is a variety of albite. It occurs in large cleavable masses, with disseminated grains
of quartz, in veins cutting Laurentian strata. The specimens exhibited were obtained from
Dr. James Wilson, of Perth, the discoverer of the mineral, who collected them in the locality
indicated. A vein of the same character occurs on the north side of Stony Lake, near the
mouth of Eel Creek in Burleigh. Its course is about N 65° E, and it intersects a white crystal-
line limestone, interstratified with blackish-grey gneiss. The vein consists of a fine grained
mixture of reddish white albite and quartz, in which are enclosed large cleavable masses
of the opalescent albite, with occasional masses of fine granular black tourmaline.—Lauren-
tian.

Orthoclase (Perthite.)

   a. Specimen cut and polished.

   This mineral, which is the perthite of Thompson, occurs in large cleavable masses, cons-
stituting, in association with quartz, a pegmatite, which occurs in considerable veins, cutting
the strata of the Laurentian series. It is a variety of orthoclase feldspar, presenting different
shades of mahogany-brown, the colors being arranged in bands. The surfaces of one of the
cleavages present golden reflections, emanating from a multitude of small points, and the
mineral very much resembles aventurine, or sunstone. These specimens were obtained
from Dr. James Wilson, the discoverer of the variety.—Laurentian.

Jasper conglomerate.

   a. Specimens intended for a vase.

   This beautiful rock consists of white quartzite, in which are imbedded a multitude of
blood-red jasper pebbles, and occurs in mountain masses in the Huronian series. While
the enclosed jasper pebbles constitute a material fit to receive the work of the jeweller,
the whole rock is capable of being applied to the manufacture of vases and such like
objects of vertu. Many boulders of the rock lie scattered along the north coast of Lake
Huron, and they are abundant at the Bruce Mines.—Huronian formation.
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Epidosite.

   a. Specimens cut and polished.
   This green rock, which is an intimate mixture of epidote and quartz, occurs in massive beds, and extends over considerable areas in the Shickshock Mountains, on the south side of the St. Lawrence, in Gaspé.—Quebec group, Lower Silurian.

10.

MISCELLANEOUS MINERALS.

Feldspar.

   a. Feldspar, from Bathurst.
   b. " " Brewer’s Mills, Rideau Canal.
   This feldspar occurs in considerable quantity on the land of Mr. Neil McEwan, and appears to form a vein of probably 20 feet in width.—Laurentian.

Sandstone for glass making.

   a. Specimen of the sandstone.
   The Potsdam sandstone in the neighbourhood of Beauharnois is in many places so free from iron as to afford an excellent material for glass making. One of the exposures giving the best examples of the stone, is at Williamstown, on the land of Mr. Donald McKillop, from which the specimen exhibited was obtained. Stone from the same formation was some years ago used for making glass at St. Johns and Vaudreuil; but it was found difficult to compete with foreign importations.—Potsdam formation, Lower Silurian.

Moulding sand.

   a. Specimen of the sand.
   This sand occurs on the surface, in patches from a few rods to several acres in extent, on the tops and sides of hills of coal sand. The best is found next the surface, and the layer seldom exceeds a foot in depth. It is the only moulding sand used in Garthshore’s extensive iron foundry in Dundas, where superior castings are made. Since to obtain a fine casting, as much depends on the quality of the sand as the skill of the moulder, the occurrence of a good quality of this material in any locality is of sufficient importance to deserve notice.—Drift.

a. Specimen of the sand.

Moulding sand occurs in two places at Owen Sound, which together may have an area of 6 acres, with an average depth of 8 or 9 inches. It is used at the iron foundries in the town, and is said to answer well.—Drift.


a. Specimen of the sand.

This is from a thin surface layer, covering between 1 and 2 acres. It is used in Cochrane's foundry in Durham, and is said to be of very good quality.—Drift.

Peat.


This peat occurs near Chambly, on the south side of the St. Lawrence, and was some years ago cut, pressed, and sold as fuel by the late Mr. Scobell. The consumption, however, was scarcely sufficient to encourage the industry. As Canada is deficient in coal, when wood becomes scarce in the progress of settlement, peat will gradually assume some importance, as a fuel in many parts of the country. Peat occurs in great abundance in many places in the province; about 100 square miles of it extend along the south front of the Island of Anticosti. Successive areas of it are met with on the south side of the St. Lawrence, from Riviere du Loup to Ste. Marie de Monoir, opposite Montreal; on the north side it occurs at La Valtrie and other places. Large peat bogs occur between the Ottawa and St. Lawrence, and there are many of the same character to the westward. The peat, which is sufficiently matted to hold together when dried, usually supports a growth of prairie grass, or ericaceous plants, or of tamarac trees. That which occurs in cedar swamps is deficient in the fibrous plants which give it cohesion, and it falls to powder when dried.—Alluvion.
DESCRIPTIVE CATALOGUE

OF A COLLECTION OF THE

CRYSTALLINE ROCKS OF CANADA.

BY T. STERRY HUNT, F.R.S.

This collection, sent by the Geological Survey of Canada, is intended to illustrate some points in the natural history of its rocks, and is divided into four parts, which are as follows:

I. Laurentian Rocks, 50 specimens, green ticket.
II. Huronian Rocks, 20 " blue "
III. Lower Silurian Rocks, 60 " yellow "
IV. Eruptive Rocks, 20 " white "

Of these, the first three are from stratified systems, and are generally distinguished as primitive or metamorphic rocks. As, however, we conceive eruptive rocks to be nothing more than displaced and altered sediments, we prefer to describe the whole collection as metamorphic or crystalline rocks, distinguishing the stratified masses which have not been displaced, as indigenous, and the eruptive ones as exotic crystalline rocks.

In the present collection, we have endeavored to do no more than present a few characteristic varieties of the principal types of rock met with in the three indigenous series. In the first and third of these, nearly all the great classes of crystalline rocks occur, and, with characteristic differences, will be found represented in each. The second series offers but a limited variety of rocks, many classes being imperfectly, or not at all represented. In the fourth division, we have selected only some of the more interesting varieties of the exotic rocks which occur in the vicinity of Montreal.
In the study of rocks, it is not possible to apply with exactness the rules of a natural-history classification, but we may conveniently arrange them in the following mineralogical groups:

1. **Silicious rocks**: as quartzite, chert, and jasper.

2. **Aluminous silicated rocks**: a, containing alumina chiefly in the form of a mineral of the feldspar family; b, as a mica or chlorite; c, as a silicate of high specific gravity, such as epidote, garnet, or chloritoid.

In this group, the feldspathic rocks are in great part reducible to two classes, 1st, **Orthosites**: in which the chief mineral is orthoclase, including trachyte, orthophyre, syenite, granite, gneiss, and argillite. 2nd. **Anorthosites**: having as their basis anorthic or triclinic feldspars. These rocks, through the introduction of hornblende, pass into diorite, and with pyroxene give rise to diabase and dolerite.

3. **Non-aluminous silicated rocks**: including serpentine, talc, pyrallolite, chrysolite, hornblende and pyroxene; the latter two minerals sometimes including a portion of alumina.

4. **Carbonated rocks**: limestone, dolomite and magnesite. These divisions suffice for our present purpose, though they exclude many substances forming rock masses, such as sea-salt, sulphate of lime, oxys, hydrates and carbonates of aluminum and iron, carbonaceous minerals, etc.

**ROCKS OF THE LAURENTIAN SYSTEM.**

The rocks of this system are the oldest known on the globe, and are widely spread in North America; where they are traced from the coast of Labrador to Lake Huron, and thence northward to the Arctic regions. Along the north side of the St. Lawrence, they form the Laurentide mountains, and in New York, to the west of Lake Champlain, the Adirondacks. The Laurentian system has been identified by Sir Roderick Murchison in the western Islands of Scotland and the adjacent coast, where it forms what was, until recently, termed the fundamental gneiss. The primitive gneiss of Scandinavia also probably belongs to the same ancient system.

The Laurentian rocks of Canada consist in great part of orthoclase gneiss, with quartzites, sometimes conglomerate, and crystalline limestones and dolomites. The total thickness of these strata is estimated at not less than 20,000 feet. Besides these, there is a great formation of anorthosite rocks, amounting to several thousand feet in thickness. These latter overlie the orthoclase and limestone series; and there are reasons for supposing a want of conformity between the two. A very distinctive and characteristic feature of the Laurentian system is the absence, so far as yet examined, of anything like argillite or clay slate. The metalliferous contents of this system are chiefly beds of magnetic and oligist iron, in the gneiss and limestone series. Iron and copper pyrites are also met with in interstratified layers, the former cobaltiferous; and both of these sulphurets, together with blende and galena, are met with in veins which cut these strata, but are as late as the Silurian period, the overlying strata of which they sometimes intersect. In the anorthosites, the only ores met with are beds of titaniferous iron or ilmenite.

2. " white, with garnets, River Rouge.

3. " pink, " River Ouitchawan.


5. " pink, granular, with garnets, Grenville.

6. " epidotic, " Carleton Place.

The most characteristic gneiss of the Laurentian series is represented by 1, which forms great mountain masses, and is so coarse grained, that, except in large masses, it might be taken for an intrusive granite. The mica which it contains, is often black, and sometimes associated with hornblende; giving rise to syenitic gneiss. Small portions of a white triclinic feldspar (albite, or oligoclase), are occasionally found with the red or pink orthoclase; and some coarse grained pegmatites, which are perhaps intrusive, consist of albite and a little quartz, with only small portions of orthoclase. The white gneiss, 2, is porphyroid, holding large cleavable masses of a pure orthoclase, in a granular mixture of the same mineral, with a little quartz and white mica, and garnets. The red gneiss, with green compact epidote, is met with in several localities in Canada. Some varieties of the Laurentian gneiss become fine grained and micaceous, passing into mica schists; but these are of comparatively small amount.


8. " pure, Rawdon.

Beds of red garnet rock are not unfrequent among the quartzose gneisses and quartzites. In the former, the mineral sometimes forms thin layers, marking the stratification. In the latter, small crystals of garnet often abound, particularly near to the limestones, and sometimes give rise to masses like 7, or to beds of a rock like 8.

9. Quartzite, conglomerate, Bastard.

9A. " Rawdon.

The above conglomerate is interstratified with white crystalline limestones, holding graphite and chondrodite. It is worthy of note, that, while some of the pebbles are of vitreous quartz, others are of sandstone, in which the layers of sedimentation are very apparent.
10. Anorthosite, granitoid, Abercrombie.


12. " " with ilmenite, Château Richer.


15. " " white, Rawdon.

16. " " granitoid, violet, River Kenogami.

17. " " greenish, ———?

18. " granular, Rawdon.


20. " " greyish, Rawdon.

21. " " bluish-green, ———?

22. " gneissoid, with ilmenite, Château Richer.

23. " " " much pyroxene, ———?

24. " " " hypersthene, Rawdon.

25. " " granitoid, with ilmenite, Château Richer.


The above seventeen specimens show the principal varieties of these remarkable rocks, which are seen, at intervals, from Lake Huron to Labrador. They often form belts of many miles in breadth, which, as before said, overlie, apparently unconformably, the orthoclase rocks and limestones. A notable feature in this formation is the almost total absence of other rocks; in some portions of their distribution these anorthosites are seen to be interstratified with thin bands of orthoclase gneiss, and more rarely with quartzite; but great masses, of thousands of feet in breadth, are found to be made up of alternating varieties of these anorthosites, which, as will be seen, exhibit great varieties of texture. Coarsely granitoid rocks abound, consisting of large cleavable masses of feldspar aggregated, or imbedded in a granular base. We find granular rocks of every grade, passing into compact and impalpable varieties with a conchoidal fracture. The composition of these feldspars varies from that of andesine, with 60 per cent. of silica (12), to varieties near anorthite, with only 47 per cent. of silica (the bytownite of Thompson). A beautiful pale blue cleavable variety (11) contains 57 per cent. of silica, being intermediate between andesine and labradorite, while many others yield from 52 to 55 per cent. The white granular rock, (15) and many others, have the composition of pure labradorite. The bases, besides aluminas, are lime and soda, with a little potash, and traces of iron and magnesia. Ten analyses of chosen specimens of these feldspars, holding from 47 to 60 per cent. of silica, gave for
CRYSTALLINE ROCKS OF CANADA.

their mean composition; silica 56.00, alumina 27.30, lime 9.42, soda 4.84, potash 0.84, magnesia, oxide of iron, water and lose 1.60 = 100.00. The oxygen ratios of the silica, the alumina, and the lime and alkalies, in the above mean, are as 7.0 : 3.0 : 0.96. From their variations in composition, we have been led to regard these triclinic feldspars, whose density ranges from 2.67 to 2.73, as indefinite crystalline mixtures of the two homoimorphic species, anorthite and albite. (See L. E. & D. Philos. Magazine for May, 1855; where also will be found analyses and descriptions of these feldspars.)

The crystalline varieties of this rock often exhibit, in great perfection, the striae resulting from their polysynthetic mazes, and are sometimes beautifully opalescent: the original Labrador feldspar is from this formation. The foreign minerals of these rocks are few in number: quartz has been seen in small portions, but is a rare accident; granular red garnet is sometimes found marking the lines of stratification, generally with pyroxene, and epidote is said to occur with the anorthosites of the Adirondacks. A brownish-black mica, probably biotite, is met with in small quantities in the granitoid varieties, but pyroxene more abundant. It is sometimes dark green and granular, occasionally predominating in small beds, so as to form a pyroxenic rock, in which small kernels, or lenticular masses of cleavable feldspar are imbedded. In other cases, where its quantity is smaller, it may be seen passing into a brownish lamellar variety, like hypersthene; the typical form of which however occurs without any association of granular pyroxene. Hypersthene is seldom an abundant mineral; it passes from brownish-black, with bronze reflections, to a clear greenish variety, like diatase. Small amounts of carbonate of lime are occasionally met with, disseminated in the granular varieties of these anorthosites. Hematite is a characteristic mineral, sometimes in thin layers, marking, with pyroxene, the sedimentary layers; at other times in larger masses, or even in beds of great size; as at Château Richer, where it is mixed with rutile.

The predominant colors of these anorthosites are various shades of blue, passing into greenish and yellowish, rarely reddish, and sometimes nearly pure white. The lustre of the cleavable feldspars is vitreous, of the granular varieties waxy or dull. The weathered surfaces are always of an opaque white; but for which, some of the white granular anorthosites might be mistaken, at first sight, for quartzites.

The nomenclature of these rocks presents some difficulties. The name of labradorite-rock sometimes given, is applicable only to certain varieties, and the same may be said of hyperstheneite and hyperite, when great masses of the rock are destitute of hypersthene. I have preferred to designate the granitoid, porphyroid, gneissoid, granular, and compact varieties of almost pure anorthic feldspar, which make up the great mass of the formation, as normal anorthosites. The interstratified beds, in which granular pyroxene predominates, would come under the denomination of dolerite or diabase, and the varieties with bronzite or diatase would, by most lithologists, be called euphotide or gabbro. Both of these names, however, would be wrongly applied; the name of gabbro, as Brongniart has shown, belongs to a diatasic ophiolite, or serpentine rock. The euphotide of Hady consists of smaragdite (a mixture of hornblende and pyroxene) with saussurite, which as we have shown, is a compact epidote or zoisite, with a specific gravity of 3.38 - 3.39. Hence neither gabbro nor euphotide are feldspathic rocks, although the euphotide of Mount Rose occasionally includes portions of a cleavable triclinic feldspar, and thus presents a transition to the diatasic variety of diabase, with which modern lithologists have confounded this epidotic rock. (See Contributions to Euphotide and Saussurite, Am. Jour. of Science, March, 1859.)

The name of diorite is, by good authorities, restricted to rocks whose dominant elements are triclinic feldspars with hornblende. In smaragdite, however, we have a mixture of hornblende with pyroxene, and in many of the so-called euphotides, according to G. Rose, the hornblende entirely replaces the pyroxene; thus forming a transition between diorite and diabase; under which latter name we propose to include the compounds of triclinic feldspars with every variety of pyroxene, except the black augite of the basalt group.
To the rocks composed of augite and triclinic feldspars, are to be reserved the name of dolerite. The pyroxenic anorthosites of the Laurentian series are then varieties of diabase, which includes hypersthenite or hyperite, the gabbro of Rose, and the euphotide of most modern authors, which latter rocks are forms of diablastic diabase, passing into diorite.

It may here be remarked that although these anorthosites, as well as the gneissic rocks of the Laurentian series, are traversed by joints in various directions, nothing corresponding to slaty cleavage has ever been remarked, and that the lamination of these masses is apparently, in every case, coincident with, and dependent upon, the original stratification of the sedimentary layers. It is only in the Huronian and Silurian series that we meet with a foliation distinct from the bedding; this is confined to the argillites, and is wanting in the more crystalline sediments.


This rock has not yet been met with in Canada, but forms considerable masses at several localities among the Laurentian series, in St. Lawrence county, New York; where it is associated with beds of oligist iron ore, and had been regarded as a serpentine, until Prof. C. U. Shepard showed its aluminous character, and gave it the name of dysyntribite. The subsequent analyses of Smith and Brush showed it to be a hydrous silicate of alumina and potash, containing, accidentally, portions of lime and magnesia, and having the composition of the agalmatolite of China. Prof. Brush has since found this mineral crystallized in hexagonal prisms, with pyroxene, at Diana. It has the composition of the mineral gieschite, with which it seems identical, and also agrees with pineite, and with what had been previously named wilsonite. This latter occurs in crystalline masses, with pyroxene and mica, in the Laurentian limestones in Canada. An analysis of the massive agalmatolite from Diana, gave to Smith and Brush, silica 46.70, alumina 31.01, peroxÿd of iron 3.69, potash, with traces of soda, 11.68, water 5.30 = 99.79. Further illustrations of this rock will be given in the descriptions of the Silurian rocks.

28. Steatite, Elzivir.

The mechanical and chemical analysis of this unctuous foliated rock, shows it to consist of tale, with small portions of quartz and of magnetic iron.

29. Pyraliolite, Grenville.

This rock forms considerable beds among the Laurentian limestones, both in Canada and New York; where it was first recognized by Prof. Emmons, and described by the name of reuselsercite. It appears, however, to be identical in chemical and physical characters with the pyraliolite of Nordenskioeld, whose name must take priority. This substance occurs massive, granular, and crystallized in the form of pyroxene. It has a specific gravity of 2.7 to 2.8, and a hardness of 2.5 to 3.5, is unctuous, greenish in color, and except in its crystalline texture, cannot be distinguished from a compact tale; with which it is identical in composition. One of several concordant analyses gave silica 61.69, magnesia 31.06, protoxyd of iron 1.35, water 5.60 = 99.79. It is in fact a monoclinic tale, and has, without any good reason, been regarded as a product of the alteration of pyroxene.
30. Pyroxene rock, with hornblende crystals, Madawaska.

31. Hornblende rock, Indian River.

This is a stratified rock, consisting of hornblende with a mixture of feldspar. It perhaps belongs to the anorthosite formation, and may be considered a diorite.

32. Pyroxene rock, with sphene, Chatham.

Small beds and interstratified masses of pyroxene rock, occur among the Laurentian limestones, and present many varieties. In 30, we have a greenish-white massive and crystallized pyroxene, with a density of 3.27; associated with crystals of a dark green aluminous hornblende (pargasite), of density 3.05, and sometimes with black tourmaline. In 32, we have a rock made up of green pyroxene, with calcareous spar and quartz; holding crystals of sphene. Sometimes a feldspar enters into the aggregate, and the rock consists of quartz and pyroxene, with an orthoclase, which is found to contain potash, with but little soda. The loxooclase of Breithaupt, which belongs to a compound rock of this kind, is, however, according to Smith and Brush, an orthoclase, in which soda predominates; and in a similar rock, with sphene, is occasionally found a triclinic species, like oligoclase. Scapolite (a dimetric feldspar) sometimes takes its place, giving rise to a pyroxene and scapolite rock.

33. Ophiolite, opaque and earthy, Calumet Island.

34. " pale green, Burgess.

35. " retinalite, Talon Portage.

36. " calcareous, yellow, Grenville.

37. " " greenish, "

We distinguish by the name of ophiolite, all rocks with a base of serpentine. In the Laurentian series, ophiolites occur, interstratified with the limestones, but offer few varieties. Their colors are usually much paler than those of the Silurian series, from which they differ in containing smaller proportions of the oxides of iron, and in the absence of those of chrome and nickel; which are constantly present in the latter. The Laurentian ophiolites are sometimes, however, of a dark red color, from the presence of disseminated peroxide of iron. The retinalite of Thompson is but a light-colored and very pure serpentine, which is noticeable for its low specific gravity, 2.36—2.32, and its large proportion of water, which equals 15.0 per cent. These ophiolites sometimes include mica; and the calcareous mixture which they hold, is often dolomitic.
38. Limestone with apatite and phlogopite, Burgess.

39. " " chrondodite and graphite, Newborough.

40. " " apatite, fluor, and spinel, Ross.

41. " " brown tourmaline, Ross.

42. " " quartz, Bastard.

43. " " with pyroxene, ——— ?

44. " " grey, with hornblende, Marmora.

45. " " graphite and sphene, Grenville.

46. " " pyroxene, Horton.

Four bands of crystalline limestone have been identified in the great Laurentian system, which are equal in volume to the ordinary limestone formations of the fossiliferous rocks. Reposing upon a vast thickness of orthoclase gneiss, we have the lowest limestone band, of about 1500 feet; to this succeed about 4000 feet of similar gneiss, followed by a second limestone formation of 2500 feet; including two bands of quartzite and hornblende orthoclase gneiss, equal to one half the volume. Following this, are 3500 feet of orthoclase gneiss, with quartzites at the base, and a third limestone band above; whose thickness varies in different parts of its outcrop, from 60 to 1500 feet. This is overlaid by 1500 feet more of gneiss, and a fourth thin band of limestone; followed by 3600 feet of quartzite and gneiss, exhibiting towards the summit, interstratified portions of anorthosites, which mark the passage to the succeeding formation. The thicknesses assigned to these masses are, however, only approximative.

The Laurentian limestones contain most of the mineral species which are met with in the crystalline limestones of other regions. Among them are apatite, fluor, wollastonite, hornblende, pyroxene, chrondodite, phlogopite, orthoclase, oligoclase, scapolite, garnet, idocrase, tourmaline, serpentine, loganite, agalmatolite, clinohumite, volcknerite, quartz, spinel, corundum, zircon, sphene, iron and copper pyrites, and graphite. Many of these minerals, such as serpentine, chrondodite, graphite, and mica, are disseminated so as to mark the stratification of the limestones. The mica, in the pure limestones, generally occurs in small scales, but sometimes in large crystals. These last are, however, most frequent in pyroxenic beds, and often with a soft stratoic mineral, having the form of pyroxene, and the composition of pyroscleite; to which it sustains the same relation as pyrallolite does to talc, and constitutes a new mineral species, called loganite. The magnesian mica, or phlogopite, often yields plates more than a foot square; which may be seen in the accompanying collection of economic minerals.

The contortions in the stratification of the limestone, show that it was once in a plastic condition, and the traces of its movement at that time, are curiously preserved, in several places, by thin interstratified layers of quartzite; which have been not only folded and broken, but twisted and rolled upon themselves, as leaves of paper would be in an agitated liquid. Occasionally we see the limestone extended among the overlying and broken layers of quartzite or of gneiss, and taking, for short distances, the form of an exotic rock.

Phosphate of lime sometimes occurs in disseminated crystals or rounded masses, in these limestones. It is a fluor-apatite, with but about one two-hundredth of chlorine, and is occasionally accompanied by fluor-spar, as in 40. These beds have been traced for several miles in the limestone, and are sometimes associated with layers of nearly pure crystalline
CRYSTALLINE ROCKS OF CANADA.

apatite. To one who is accustomed to look upon the graphite, and the great beds of iron ore in this system, as evidences of the intervention of organic life during the Laurentian period, these layers of phosphate of lime seem to be accumulations of coprolitic matters, from the animals (perhaps marine) of that period; in fact, the ancient representatives of modern guano beds. In the unaltered strata at the base of the Silurian system, layers of both limestone and sandstone, abound in phosphatic coprolites, apparently derived from the Lingulæ, Orbiculæ, Conulariæ, and Sertulæs, of those early times; the shells of all of which have been shown by us to have essentially the composition of the skeletons of vertebrate animals.—Am. Jour. of Science (2), xvii. 236.

47. Dolomite, with green mica, Indian River.
48. " " white mica, Madawaska.
49. " " tremolite, "

Great beds of crystalline dolomite, and of limestones more or less magnesian, occur, interstratified with the purer limestones of this series. They are often very fine grained, and sometimes resemble statuary marble; others contain a portion of peroxide of iron, and weather to a reddish-brown. Foreign minerals are less abundant in the dolomites than in the limestones; but besides mica, tremolite, and quartz, serpentine sometimes abounds, forming a dolomitic ophiolite. We shall consider the chemical and geological relations of dolomites, and the theory of their formation, in describing the Silurian rocks.

II. ROCKS OF THE HURONIAN SERIES.

The rocks which have been designated as the Huronian series, rest upon those of the Laurentian system, and are in part made up of the ruins of the latter. The unaltered and horizontal Lower Silurian strata, in their turn, repose upon the inclined and metamorphosed Huronian rocks, which are therefore regarded as constituting a distinct and intermediate formation. This seems, from its geological horizon, not less than from its lithological characters, to correspond to the quartzose division of the Primitive Slate Formation of Scandinavia. The Huronian series is met with at Lake Temiscaming, on the Ottawa, and on Lakes Huron and Superior. It is not known farther eastward, but it is not unlikely that it constitutes some portions of the Azoic rocks of the Upper Mississippi, and of Arkansas and Missouri.

The thickness of the Huronian series on the north shore of Lake Huron, is approximately estimated at 18,000 feet. Of this, more than 10,000 feet are quartzites, which are sometimes schistose and micaceous. The remainder consists of chloritic and argillaceous slates, which occasionally hold epidote, and, like the quartzites, often become conglomerates. Three small bands of impure limestone occur in this formation, two of which are associated with layers of chert or hornstone. Throughout the whole formation, are interstratified great beds of crystalline gneiss or diorite, sometimes several hundred feet in thickness.

We remark in this series of rocks, but a small amount of carbonate of lime, and an absence of well characterized gneiss or orthoclase feldspar rocks. An impure ferruginous serpentine has been observed in the series, near Marquette, but no steatites nor talc slates. Its metalliciferous minerals consist of beds of specular iron, to which species the great mines of Marquette, in Northern Michigan belong, and of large quantities of sulphurates of copper. The copper ores sometimes occur disseminated in the diorites or chloritic slates, but more generally in well-defined veins of quartz, which traverse the dioritic rocks.
DESCRIPTIVE CATALOGUE.

1. Quartzite, white granular, Island near Grant's Island.
2. " " vitreous, Grande Batture Point.
4. " " schistose, Lacloche.
5. Limestone, slaty, Clear Lake.
6. Quartzite, conglomerate, with jasper pebbles, Bruce Mines.
7. " " " "

Quartzite may be said to be the predominant rock in the Huronian series. Its colors are white, gray, brownish, and sometimes greenish or reddish, and its texture is various; it being sometimes vitreous, and at other times, a granular sandstone. It is not unfrequently schistose, and sometimes slightly micaceous or feldspathic, but true gneiss and mica slate have not been met with in this series. These quartzites often become conglomerate, from the presence of various colored pebbles of quartz and jasper. The latter are frequently blood-red in color, and being imbedded in a white or a greenish base, constitute a very beautiful rock.

8. Argillite, bluish, talcoid, Spanish River.
9. Hornstone, in limestone, Chert Point, Lake Superior.
10. Limestone,
11. " Lake Huron.

The limestones of this series are but small in amount. One band of 300 feet in thickness has however been traced for considerable distances. Its colors are chiefly greyish, greenish, or buff, rarely white, and its fracture is conchoidal, and sometimes granular. It is often ferruginous and yellow-weathering, and is somewhat magnesian. Thin silicious layers give to its weathered surface a very uneven aspect. It is strikingly contrasted with the Laurentian limestones, by the absence of any pure crystalline varieties, or imbedded crystalline minerals. Two other bands, of 200 and 400 feet respectively, consist of similar impure limestones, with regular layers of yellowish chert, the latter predominating. Beds of this chert or hornstone are sometimes interstratified with the adjacent quartzites.

12. Argillite, greenish, Grant's Island, Lake Huron.

Beds of clay slates are sometimes met with in this series; they are occasionally bluish and talcose in their aspect, and at other times greenish, and apparently somewhat chloritic. We have noted the absence of clay slates from the Laurentian system; and their presence in the Huronian series, shows a condition of things approaching to that of the Silurian period, when we find these rocks in much greater abundance.

15. " " Clear Lake.

16. " " conglomerate, Echo Lake.

17. " "

Great masses of a greenish slaty rock are met with in this series, which varies in hardness and texture, from a silicious slate, passing into hornstone, on the one hand, to an argillaceous or a chloritic slate, which is sometimes epidotic, on the other. These slates frequently include pebbles of crystalline rocks, which are chiefly feldspathic, and derived from the Laurentian strata. With these are, however, sometimes mingled others of quartz and of various colored jaspers. The pebbles vary much in their amount, and the rocks pass from ordinary slates, to what have been designated in the descriptions of this series, as slate conglomerates. The matrix of these is sometimes an argillaceous or chloritic slate, and occasionally becomes very quartzose, passing into a quartzite; so that it is not easy to draw the distinction between the conglomerate slates, and the jasper conglomerates of the quartzites.

18. Diorite, compact.

19. " fine grained, Dirty Lake.

20. " coarse grained, "

The diorites or greenstones of the Huronian series are intercalated in beds, alike with the quartzite and the argillaceous and chloritic members. They are sometimes coarse grained and crystalline, being made of dark green hornblende and a greenish feldspar. In other parts, the rock becomes fine and even compact in its texture, and it is frequently porphyritic from the presence of crystals of feldspar. Great masses of the rock become schistose, and are often intermingled with a considerable amount of chlorite, passing into dioritic and chloritic slates; which are often associated with a considerable amount of epidote, generally granular or imperfectly crystallized. In one locality, amygdaloidal strata, holding in their cells, quartz and calcite, are found interstratified with the chloritic and the porphyritic beds. In some few instances, the feldspar in the coarse-grained diorite becomes reddish, and the rock includes a little quartz, passing into a variety of syenite. The Huronian series is traversed, like the Laurentian, by dykes of greenstone trap; but the great beds of diorite just noticed, are considered to be altered sedimentary rocks.
III. ROCKS OF THE SILURIAN SERIES.

The Notre-Dame and Shickshock Mountains are the north-eastern prolongation of the great Appalachian chain, which extends from the Gulf of St. Lawrence, nearly to the Gulf of Mexico. These mountains, at least in Canada and New England, are altered sediments of Paleozoic age, and are referred to the Quebec group; which corresponds to the inferior part of the Lower Silurian series. They attain, in some places, a height of more than 4000 feet above the sea, and appear to be generally synclinal in their structure. The rocks are highly metamorphosed in the mountainous region, which constitutes a narrow belt, but on the north and west of this are found in a comparatively unaltered state. These hills, and the region around them, offer almost every variety of metamorphic sediments, but they are very deficient in intrusive rocks, of which scarcely a single dyke can be met with. The country on both sides of the altered mountainous belt, abounds in intrusive masses of various kinds, some of which will be described in the succeeding portion of this catalogue.


1a. Gneiss, granitic, St. Joseph.

Great masses of orthoclase gneiss are met with in this series. They are generally fine-grained, and are more quartzose than those of the Laurentian system; with which the practiced observer will never confound them. The coarse-grained and porphyritic reddish and white varieties are never met with, and the gneiss is generally of pale greyish or greenish hues. In some cases, great portions of it are so destitute of marks of stratification, that but for their relations to the adjacent beds, they might be taken for intrusive masses. The mica is generally white or greyish, and in small quantity.

2. Anorthosite, Melbourne.

2a. " Orford.

3. " with serpentine.

Rocks composed of triclinic feldspars, and representing the anorthosites of the Laurentian system, are common in this series; they are however never coarsely crystalline, and are often compact. In some cases the feldspar approaches to albite or to oligoclase in composition. Through an intermixture of hornblende, these rocks pass into diorite.
4. Diorite, St. Francis.

5. " Tring.


7. " "


In the diorites of this series, the feldspar is sometimes the predominant element. One from Orford was found, by analysis, to consist of sixty-four parts of albite, and thirty-six of hornblende; another contained seventy-four parts of a feldspar, which was near albite in composition, but contained as much potash as soda. Others of these diorites exhibit a predominant of hornblende, often mingled with a chloritic mineral, and constitute veritable greenstones; which, however, appear to be in all cases sedimentary rocks. They are frequently so finely granular as to appear at first sight homogeneous, while at others they are rather coarsely crystalline, or sometimes porphyritic, from the presence of large feldspar crystals. The specimen from St. Joseph is associated with compact white garnet and crystallized hornblende.

9. Epidosite (epidote and quartz), Melbourne.

10. " schistose, with oligist iron, "

11. " chloritic, with epidote nodules, St. Armand.

12. Epidotic rock, with calcite and argillite, St. Joseph.

Epidote is a characteristic mineral of great portions of this series. Sometimes it forms with quartz, a fine-grained compact rock, which is found in thick beds in the Shickshock Mountains. At others, the epidote is disseminated in nodules, in a fine grained silicious rock, which often becomes chloritic or argillaceous.


A massive white lime-alumina garnet occurs in beds in this series, sometimes in contact with ophiolite, or mingled with feldspar and hornblende (as in 8), or with an admixture of serpentine. This garnet-rock is extremely tough, in some cases imperfectly crystalline, and has a specific gravity, when pure, of 3.53. Other specimens, probably mingled with feldspar or hornblende, and greyish or greenish in color, have a density of 3.29-3.4.


15. " " " "

These specimens, which should be placed with 12, from the same locality, are from a great mass of argillaceous rock, which passes into red shale in some parts, and in others, is concretionary in its structure. It would appear as if the clay had originally contained septaria, the fissures in which, as well as the inter-tices, have become filled with epidote, which is often crystallized, calcite, quartz, and sometimes talc. These altered argillites are in the immediate vicinity of the ophiolites, and, in some specimens, much resemble the gabbro rossi of the Italian geologists.

This soft grey schistose rock, a bed of which has been wrought as a variety of potstone, has nearly the composition of a hydrous mica, with only three per cent of alkalis, and fifty-one per cent of silica.


These mica-schists are very variable in their nature, and often highly quartzose; not unfrequently they have the aspect of what are called talcose slates, without however containing any magnesia, and owe their peculiar characters to a mica like that of 16, or perhaps to pholerite or pyrophyllite. Pholerite is sometimes found in a pure state, in fissures in the sandstones of this series; and pyrophyllite forms beds, resembling steatite, in the same formation in the southern United States; where it also occurs crystallized with quartz.


23. Mica-schist.


26. “ bluish “

27. “ with orthoclase and quartz, Cleveland.


29. “ with red orthoclase, Cleveland.

The argillaceous rocks of this series present many varieties, from roofing-slates, and talcoid and plumbaginous shales, to others which are more or less chloritic or micaceous. The specimen 27 is remarkable from containing small oval masses of regular outline, consisting of orthoclase and quartz. Their exterior portion is generally of feldspar, the centre being filled with quartz; but sometimes the one or the other is wanting, and the kernels consist of quartz or of feldspar only. These oval masses, which are from one-eighth to one-half an inch in length, have their greater diameters parallel. The rock might be called an amygdaloid. Some portions of these argillites are penetrated by small veins and irregular masses of bright red orthoclase. This feldspar is occasionally found in veins with quartz, chlorite, and bitter-spar, intersecting these slates.

Chloritoid is abundant in quartzose mica schists, in this series. It is generally in small plates, but sometimes in tables one-fourth of an inch in diameter, often arranged in spherical aggregations. It has a specific gravity of 3.5, and the usual composition of the species. Chloritoid is identical with the barytophyllite and the sismondine of different authors. It is also supposed to be the phyllite of Thompson, and the osltreite of Haüy, both of which closely resemble it in appearance.


32. “ “ “


Great beds of a rock made of scales of specular iron, with quartz and chlorite, are met with in the altered Silurian strata. They are sometimes rich iron ores, and at other times contain but small portions of the metallic oxide. These specular schists often include a portion of titanic acid, which is occasionally seen in the form of rutile or of sphene, crystallized in veins, sometimes with feldspar. These rocks are apparently identical with the itabirite of Brazil.

34. Magnetic iron in dolomite, Sutton.

Magnetic iron ore is often found in these rocks, in irregular beds or masses in serpentine. In Vaudreuil, there is found a bed of granular ore, of which two-thirds are pure magnetite, and the remainder ilmenite; the two being intimately mixed. Grains and octahedral crystals of magnetic also occur in the chloritic schists, and in the present specimen, the crystals are so abundantly disseminated in some parts of a bed of chloritic dolomite as to constitute a valuable iron ore. This dolomite is remarkable for containing eight per cent. of carbonate of manganese.

35. Copper pyrites in chloritic limestone, Ascot.

36. Variegated copper in micaceous schist, Sutton.

37. “ “ “ “

Copper is abundantly distributed in this formation, generally disseminated in the beds, and forming an integral portion of the rock, in the shape of grains or lenticular patches. The yellow and variegated sulphurites, copper glance, and sometimes native copper, are met with alike in quartzose, argillaceous, micaceous, and chloritic slates, in limestones, and in dolomites. At Acton, the latter two ores form the cementing material of a conglomerate rock, made up of limestone and siliceous matters. The copper in these strata seems to have been a contemporaneous deposit from aqueous solutions.
38. Hornblende rock, with garnets, Shickshock Mts.

Beds of black crystalline hornblende rock, including small crystals of red garnet, occur with the serpentines of Mount Albert. In many other parts, hornblende in the form of actinolite, or a tough, fibrous variety allied to it, forms beds of great thickness.

39. Diabase rock, Orford.

Diabase is abundant, not only as a component of some ophiolites, but sometimes forming a rock, either by itself, or with a little mixture of an amorphous mineral, which approaches to pyroscelite in its composition.

40. Ophiolite, (serpentine,) Orford.

Under the name of ophiolite we include those rocks which have serpentine for their base. The normal ophiolites are nearly pure serpentine, while some are mixtures of serpentine and carbonate of lime (calcareous ophiolites), and others dolomitic and magnesitic ophiolites; containing respectively dolomite and carbonate of magnesia, often in large proportions. All of these varieties are met with in Canada, or in the adjacent state of Vermont. These compound ophiolites are sometimes porphyritic from the presence of diabase (the Italian gabbro). At other times, they have the aspect of conglomerates, exhibiting rounded or angular masses of pure serpentine of various sizes, imbedded in a dolomitic paste, itself more or less colored by intermingled serpentine. A magnesitic ophiolite from Vermont has a gneissoid structure, due to the arrangement of the crystalline magnesite spar, with lamellae of talc, apparently marking planes of stratification. The ophiolite of Mount Albert is marked with red and green bands, (see specimen 59,) which have the aspect of sedimentary layers; and the relations of the ophiolite throughout this series, where its outcrop has been followed for hundreds of miles, are always those of an interstratified deposit, and never of an eruptive rock. It occurs with dolomite, magnesite, steatite, diorite and argillite, with each one of which it has been found in contact, and it seems sometimes to replace the other magnesian rocks. Its beds vary from a few yards to several hundred feet in thickness. The colors of these ophiolites are of various shades of green; generally much darker than those of the Laurentian series. A red color sometimes occurs in patches and bands, or pervades the whole mass; this, in some cases, at least, is due to an intermixture of red hematite. Foliated and fibrous varieties (baltimore and chrysotile) are frequently found in veins in these ophiolites. Chromic iron is also a characteristic mineral, in grains, or in interstratified beds or laminar masses, often of large size. Magnetic iron occurs in these ophiolites, both in grains and beds, sometimes with limesite.
The analyses of the serpentines of these ophiolites show them to contain from seven to ten per cent. of protoxide of iron, to which they owe their color, besides small portions of oxides of chrome and nickel. These two metals often occur in the magnesian rocks of this series, in the form of chromic iron and sulphuret of nickel; but are in many cases present as integral portions of the silicate. This is true, not only of the serpentines, but of the diabase and actinolite rocks, and many of the dolomites and magnesites. It would seem that chrome and nickel were constant accompaniments of the magnesian deposits of the present series. We have also detected these metals in the ophiolites of California, of Portland in Scotland, Cornwall, the Vosges Mountains, Mount Rosa and Corsica; while they are wanting in the Laurentian ophiolites of Canada, and in specimens of serpentine from Norway, supposed to be of the same formation.

45. Steatite, Bolton.

46. " with bitter-spar, Ireland.

Besides the so-called talcose slates of this series, which are for the greater part aluminous, true talc slates, or schistose varieties of steatite are not unfrequent. These are sometimes nearly pure talc, and at others mingled with hornblende, in the form of actinolite, or with bitter-spar. They yield to analysis a few thousandths of oxide of nickel.

47. Chlorite (potstone,) Bolton.

The chloritic slates of this series are often mingled with quartz and with epidote, and sometimes with specular iron. In other cases, however, beds of pure, compact, and somewhat schistose chlorite, occur.


49. " Bolton.

Magnesite rocks have been met with in three localities in this series. That of Sutton occurs with dolomite and steatite, and consists of carbonate of magnesia with some carbonate of iron, intermixed with grains of a feldspathic mineral, and a green, chromiferous mica. The magnesite of Bolton forms an immense bed, between steatite and diorite, and contains a mixture of grains of quartz, besides small portions of both chrome and nickel. In a third locality, the magnesite, which is compact, earthy, and yellow-weathering, is interstratified with argillite, and resembles in appearance many of the magnesian limestones of the region.
Dolomites, or magnesian limestones, are abundant in this series, and frequently accompany the ophiolites, into the composition of which, as we have seen, they often enter. These dolomites are generally ferruginous, often containing eight or ten per cent. of carbonate of iron, and sometimes as much carbonate of manganese. They are often mingled with a portion of clay, or of silicious sand, sometimes considerable in amount, and very frequently become conglomerates, enclosing pebbles or rounded masses of pure limestone, and more rarely of sandstone, shale, or dolomite, in a paste of ferruginous red-weathering magnesian limestone. In some cases, these rocks have the composition of a true dolomite, in which the oxides of iron and manganese replace a portion of magnesia. In others, the quantity of lime is not equivalent to the other protoxyd bases, and we have a passage to the magnesites already described; which are rocks consisting of carbonates of magnesia and iron, with little or no carbonate of lime. The carbonate of iron occasionally predominates in these, and in some instances, a bed of spathic carbonate of iron occurs. The foreign minerals of these rocks are few in number; chlorite, talc, hornblende, pyroxene, and brown garnet are sometimes met with, and a green chromiferous mica, probably allied to fuchsite, occurs, in small scales, both in the magnesites and in the dolomites. An emerald-green garnet with six per cent. of chrome, is also, in one place, associated with these magnesian rocks. With the ferruginous dolomites, are often interstratified beds of pure limestone, which frequently enclose concretionary fibrous masses, made up of concentric layers, like the recent deposits of travertine from calcareous waters.

The conditions under which these dolomites and pure limestones are associated, are such as to leave no doubt that they have been contemporaneous deposits, and to forbid the notion of the formation of dolomite by any subsequent alteration of the limestones. In a series of investigations published in the Reports of the Geological Survey for 1857 and 1858, we have endeavored to explain the origin of these carbonates of lime and magnesia, and their associations in nature. It has there been shown that when waters holding bicarbonate of soda in solution, act upon sea-water, containing chlorids of calcium and magnesium, the whole of the lime is first precipitated in the form of carbonate, with but a very small proportion of magnesia. A farther addition of the alkaline carbonate, if fresh supplies of lime salts are excluded, gives rise to a very soluble bicarbonate of magnesia, which, by evaporation, is separated as a hydrous carbonate. This, when heated alone to 800° F., under pressure, to prevent the loss of carbonic acid, is changed into magnesite, but if mingled with carbonate of lime, a double salt results, which is dolomite. The sources of the alkaline carbonate are to be found in decomposing feldspars; the surface waters from regions of feldspathic rocks, and the springs which traverse the debris of such rocks, are still, at the present day, impregnated with carbonates of soda and lime; in the latter case, they are often accompanied with oxys of iron and with rarer metals. In this way the metalliferous character of many dolomitic formations is explained. The carbonated rocks have thus been formed by a series of decompositions, the results of which are represented by the clays and argillites (which are feldspars that have lost a portion of their alkali), by the limestones and dolomites, and by the chloride of sodium in the sea, and in the rocky strata. All limestones, as well as dolomites, are the result of this chemical process, which furnishes the elements for the limestones of organic origin. Great masses
CRYSTALLINE ROCKS OF CANADA.

of carbonate of lime, in various formations, as for example the statuary marbles of Lower Silurian age, in Vermont, are purely chemical in their origin, and do not result from the metamorphism of fossiliferous limestones.

These views were first enunciated in the reports of the Canada Geological Survey, already cited, and in the Am. Jour. Science, May 1858, xxv, 302, and (Quar. Jour. Geol. Soc. London, for 1859, p. 492. In a sealed packet deposited by Cordier, with the French Academy, some years ago, and opened since his death, the same views are suggested.


The magnesian limestones, commonly associated with beds or masses of gypsum, appear to have been formed by a reaction pointed out in the above Reports; in virtue of which, solutions of bicarbonate of lime, when mingled with evaporating waters holding sulphate of magnesia, give rise to sulphate of lime, which is first separated, and to a more soluble bicarbonate of magnesia, which is deposited by further evaporation, mingled with a further portion of carbonate of lime. The sulphate of magnesia, which, in Canada, as elsewhere, often exudes from these dolomites, appears not to be due, to a subsequent reaction between the dolomite and the gypsum, but to have been an original element of these rocks.

55. Chert, Cape Rouge.

56. Sandstone, St. Nicholas.

57.  "  "

58. Agalmatolite, "

The agalmatolite of St. Nicholas, which had at first been taken for serpentine, was described, with analyses, in the Report of the Survey for 1850, under the name of parophite. The subsequent analysis of the dysuninite of Shepard, from the Laurentian series, showed the identity of the two rocks which have, as already remarked on page 67, the composition of agalmatolite or of the onkios of Kobell. The specimens from St. Nicholas form thin layers, often concretionary, in an earthy shale, which has apparently the same composition. In other localities in this series, however, the agalmatolite appears as a soft, unctuous, translucent, yellowish-green rock, which is either granular, or has an indistinctly ligneous structure, with a satiny lustre.

 Deposits of silica, which are evidently of chemical origin, and which assume the forms of hornstone or Jasper, as they include more or less argillaceous or ferruginous matter, are not unfrequent among the mechanical sediments of this series. The two specimens of sandstone from the unaltered strata of the Quebec group at St. Nicholas, are supposed to represent the granitic gneiss of the altered portions of the same formation. The cement, in some of these sandstones, is a feldspathic matter, rich in potash; and the analysis of the rock, as a whole, gives a composition identical with the mixture of quartz, orthoclase, and mica, which constitutes this gneiss. The metamorphism of these alumino-rous rocks consists, then, simply in the crystallization of the silicates of alumina and alkali in the sediments, a reaction which has taken place at no very elevated temperature; the alkaline silicates and carbonates, by which the waters of these sediments are impregnated, aiding the process. At the same time, the reactions between the silicious and argillaceous matters, and the earthy carbonates, in the presence of these alkaline solutions, give rise to chlorite, garnet, and epidote. These views, together with various experiments on the artificial formation of silicates, were published in the Am. Jour. Science for May 1857, p. 438, and the Proc. Royal Society for May 7, 1857. They are also given in the Report of the Geological Survey of Canada for 1856, p. 479; all of which appeared anterior to the first publication of Daubreé; who, in November 1857, brought forward some striking experiments in support of the theory of the metamorphism of sediments, at comparatively low temperatures, by the intervention of alkaline salts.
In the Report for 1858, p. 188, will be found some account of the results of local metamorphism of limestone near a trap dyke at Montreal. The limestone here contains a portion of an argillaceous matter, with 70 per cent. of silica, consisting of finely divided orthoclase and quartz. Where the beds have been rendered crystalline, near the intrusive rock, those substances have become saturated with lime, magnesia, and oxide of iron; and there results a silicate of these bases, with alumina, containing only 40 per cent. of silica. By similar reactions, the various silicates of lime and magnesia, both hydrated and anhydrous, may be formed; including both serpentines and talc. Steatite is however doubtless but the result of the molecular metamorphism of sepiolite, a silicate of magnesia which occurs in beds in many Tertiary deposits; and ophiolites have probably originated in beds of a similar magnesian silicate. The source of these silicates may be traced to the spontaneous evaporation of natural waters, many of which deposit silicates of lime, magnesia, and oxide of iron. The proportion of silica in solution in the waters of the Ottawa River, is one third of all the solid matters (which amount to 6 parts in 100,000), and a part of this remains dissolved, together with lime and carbonate of soda, in the concentrated water; which, like that of the St. Lawrence, deposits an earthy silicate by further evaporation. (Report of Geol. Survey for 1855-56, p. 360.)

The problem of rock metamorphism is the conversion of mechanical or chemical sediments into definite mineral species, by molecular changes, or by chemical reactions between their elements. Pseudomorphism, which is the change of one mineral species into another, by the introduction, or the elimination of some element, presupposes metamorphism; since only the definite mineral species of metamorphic rocks can be the subjects of this process. To confound metamorphism with pseudomorphism, as some have done, is therefore an error. It may be further remarked, that, although certain pseudomorphic changes may take place in some mineral species, in veins, and near to the surface, the alteration of great masses of silicated rocks by such a process, is an unproved hypothesis.

IV. INTRUSIVE ROCKS.

The results of recent geological investigations in various parts of the world, lead to the conclusion that many rocks, formerly regarded as intrusive or exotic, are really sediments, altered in situ, or indigenous rocks. Such is the case with many granites, syenites, greenstones, amygdaloids, porphyries and serpentines; all of which are represented among the altered strata of Canada. These sediments at the time of their metamorphism, were however in such a plastic state, that they were sometimes displaced and forced among the overlying and disrupted strata. It is not improbable that the intrusive granites, which are so abundant among the Devonian rocks to the south and west of the Notre-Dame Mountains, are the equivalents of the feldspathic sandstone and granitoid gneiss of the lower Silurian series. It is worthy of note, that intrusive masses are extremely rare in the Laurentian system, so far as known, except in one small area in the counties of Grenville and Argenteuil, where a succession of eruptions of dolerite, syenite, and quartziferous porphyry, occurred before the commencement of the Silurian period. In the same way, the great masses of the Lower Silurian mountains are free from intrusive rocks. To the south-east of them, however, occur the Devonian granites just mentioned, and to the north-west, along the valleys of the St. Lawrence and Lake Champlain, are a series of intrusive dolerites, diorites, and trachytes. The most remarkable of these, in Canada form a line of isolated hills, eight in number, extending about ninety miles along the line of an undulation, which runs nearly east and west, or almost transverse to the Notre-Dame Mountains, and has disturbed the Lower Silurian strata. These hills, beginning from the west, are Rigaud, Mount Royal, Montarville, Belœil, Rougemont, Yamaska, Brome and Shefford Mountains, to which may be added Mount Johnson, or Monnoir, a little to the south of this line. Brome and Shefford are on the confines of the metamorphic region. These masses, which were intruded among the members of the Lower Silurian series, have
been left by denudation, as hills, covering areas of several miles, and sometimes more than 1000 feet in height, and present great varieties in composition. Brome and Shefford are granitoid trachytes, Yamaska, partly trachyte and partly diorite; to which latter rock also belongs Boisbriand, so far as examined, and Monnoir. Rougemont, Montarville, and Mount Royal are dolerites, and Rigaud is, in great part, a granitoid trachyte. Dykes of numerous varieties of trachyte and of phonolite, cut the dolerites of Mount Royal, and the shales of the Hudson River formation. The conglomerate of St. Helen, which overlies and encloses masses of Upper Silurian limestone, as well as fragments of granitoid dolerite, is in its turn traversed by dykes of a newer rock, which is also a dolerite. The strata in the vicinity of these intrusive masses are not altered, except near the line of contact. (See page 89.) The present collection includes only a few of the more characteristic varieties of these intrusive rocks.

1. Quartziferous Porphyry, Grenville.

In the county of Grenville, the Laurentian limestones and gneisses are successively cut by intrusive masses of dolerite, syenite, and quartziferous porphyry, all of which rocks are older than the Silurian period. The last of these, which is an orthophyre or felsite porphyry, has a compact, apparently homogeneous base, inclosing crystals of orthoclase, and more rarely, grains of quartz. The color of the crystals is of different shades of red, while the base varies from black to purplish and greenish hues, and is found by analysis to consist of an intimate mixture of orthoclase and quartz, colored apparently by oxyd of iron. This porphyry receives a fine polish, and some varieties of it are very beautiful.

2. Trachyte, granitoid, with hornblende, Shefford Mountain.

3. " " " mica, Brome Mountain.

4. " " " Yamaska "

5. " compact, with pyrites, Montreal.

6. " " "

7. " " "

8. " " red-weathering, "


The mountains of Shefford and Brome are masses of intrusive rock, which break through the shales of the Quebec group; the latter, which is the larger, occupying an area of about twenty square miles. These mountains are composed of a granular rock, which might be mistaken for granite, but for the absence of quartz. It is an aggregate of crystalline grains
of orthoclase feldspar, with a small admixture of hornblende or black mica, which appare in different parts to replace one another. The rock is sometimes fine-grained, but in other parts consists of cleavable forms of orthoclase, which are occasionally half an inch in length. Small grains of magnetite, and of yellow sphenite are also sparingly disseminated. This rock, from the absence of any mineral as a cementing medium between the grains of feldspar, is very friable, and rapidly disintegrates at the surface. Its structure and composition are such that it may be designated a granitoid trachyte. The feldspar has specific gravity of 2.56. One of several concordant analyses, from different localities, gives its composition: silica 65.15, alumina 20.55, lime 0.73, potash 6.39, soda 6.67, volatile 0.50 = 99.99.

A variety of this trachyte, from a dyke near Chambly, consists of large well-defined orthoclase crystals, in a fine-grained, lamellar base, both having nearly the same composition as that just given. The vicinity of Montreal abounds in trachytic dykes, which are generally fine-grained; they are sometimes crystalline, and at others earthy in texture, and are occasionally porphyritic from the presence of feldspar crystals. They are generally white or grey, and more rarely lavender-colored or purplish in hue. These trachytes often contain disseminated earthy carbonates, in some cases amounting to from seven to fifteen per cent., and consisting of carbonate of lime, with considerable proportions of carbonate of magnesia and protoxyd of iron. These varieties of trachytes are often grey, granular and semi-vitreous, but effercesce freely with acids. The more earthy of them are sometimes weathered to a little depth, and reddish from the peroxydation of the iron. The insoluble residue of all these rocks approaches in composition to the orthoclase above described. In some cases, these trachytes contain an admixture of a hydrated silicate which gelatinizes with acids, and has the composition of a zeolite: through this admixture they pass into phonolite.


This rock forms a large dyke, traversing the shales of the Utica formation. It is a few colored compact mass, with a somewhat schistose fracture, and has a specific gravity of 2.41. It effercesce slightly, and gelatinizes with acids, and is found by analysis to consist of from forty-five to fifty-five per cent. of an insoluble potash feldspar, near to orthoclase in composition, with from thirty-six to forty-six per cent. of a soluble hydrous silicate of alumina and soda, closely approaching to natrolite; besides about seven per cent. carbonates of lime and protoxyd of iron, in nearly equal proportions.

12. Dolerite (Oligoclasic), Mount Johnson.

The isolated Mount Johnson, or Monnoir, as it is sometimes called, consists of a granul diorite, made up of black crystalline hornblende and white cleavable feldspar, with some crystals of amber-yellow sphenite. The rock is sometimes finely granular, but more generally coarsely granitoid or porphyritic; the crystals of feldspar, which is the predomina mineral, being frequently an inch or more in length. They have a specific gravity 2.63-2.65, and the composition of oligoclase. Its analysis gave silica 62.65, alumina 22.4, protoxyd of iron 0.75, lime 2.96, potash 1.80, soda 7.35, volatile 0.86 = 99.99.

The diorite of Yamaska much resembles the last, being made up of black hornblende, with a white feldspar, and small grains of sphene and magnetic iron. It is sometimes granular, but the feldspar often presents striated cleavage planes, half an inch in breadth, which have a specific gravity of 2.75 - 2.76, and a composition near that of anorthite. Its analysis gave silica 46.90, alumina 31.10, peroxyd of iron 1.35, lime 16.07, magnesia 0.65, potash 0.58, soda 1.77, volatile 1.00 = 99.42.

This beautiful diorite makes up a large part of the mass of Yamaska mountain, but the remainder is a granitoid trachyte (4). This is more micaceous than that of Brome, and consists in great part of a feldspar, which approaches oligoclase or andesine in composition.

14. Dolerite, Montarville.

15. " "


17. " " "

18. " (Peridotite), Rougemont.

19. " " Montarville.

20. " " Vermont.

The dolerites which form the mountain masses of Rougemont, Montarville, and Mount Royal, present great varieties in their composition. Some parts of the latter mountain consist of a granitoid aggregate of a greenish-white feldspar, having the composition of labradorite, with black augite. This latter sometimes prevails, to the almost complete exclusion of the feldspar, forming a crystalline augite rock. In other parts, the black and more augitic portions are arranged in short irregular bands, with a lighter and more feldspathic dolerite, as if two plastic masses, holding different proportions of augite, had been partially mingled in flowing. Grains of olivine sometimes occur in the more feldspathic portion of Mount Royal, and are still more abundant in a similar rock from Rougemont and from Montarville. In both of these masses, more or less augitic varieties occur, as at Mount Royal. The chrysolite or olivine, which is rare in the greater part of Montarville, predominates in one portion, which is a granitoid aggregate of feldspar and augite; the latter often in well defined crystals, with a little brown mica, and grains or imperfect crystals of yellowish olivine. This, in some specimens, equals forty-five per cent. of the rock, and consists of silica 37.17, magnesia 28.68, protoxyd of iron 22.54 = 99.39.

This peculiar rock, which, from the predominance of olivine or peridot, might well be separated, from dolerite, may be distinguished by Cordier's name of peridotite. It is the more worthy of attention, from the fact that olivine has hitherto been regarded as characteristic only of fine grained dolerites or basalts. As an example of an extremely coarse grained or granitoid peridotite, the specimen 20 is subjoined. This rock, which consists of great crystals of cleavable feldspar, with masses of granular chrysolite, and small portions of green pyroxene, was found in a boulder, in the Connecticut valley.
ADDITIONS AND CORRECTIONS.


" 10. Wellington Mine: for "Plans of the mine by Mr. C. H. Davie," read "by Mr Plummer."

" 13. Yale's Mine: after the second line, insert "b, Plan of the mine, by Messrs Willson & Robb." In the sixth line, for "from three to twelve inches;" read "from six to thirty inches."

" 14. St. Francis Mine: 7th line, for "The bed is about eighteen inches wide," read "The bed has an average thickness of three feet." 9th line, for "Two small excavations," read "Five or six small excavations."


" 49. Gypsum: add, "The total amount of gypsum mined annually on the Grand River, is about 14,000 tons."
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